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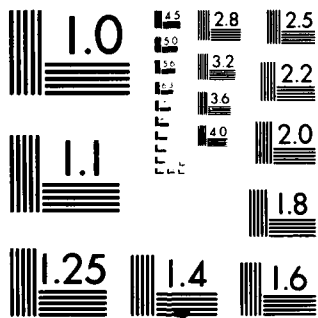
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RADAR TRACKING OF BARIUM ION CLOUDS: RESULTS OF THE PLACES EXPERIMENT

Victor H. Gonzalez

SRI International
333 Ravenswood Avenue
Menlo Park, California 94025

1 August 1981

Final Report for Period 1 September 1979—1 May 1981

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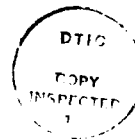
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I INTRODUCTION

The PLACES communication experiment consisted of a series of high-altitude barium releases that took place at Eglin Air Force Base, Florida. During these experiments, the AN/FPS-85 UHF radar, operating in an incoherent scattering mode, tracked the ionized barium cloud and recorded the measurements that were taken. This mode of measuring electron densities consists of receiving the aggregate reflections from individual electrons that populate the ionized region of interest. The resolution of the radar beam is sufficiently adequate to attain a spatial, as well as temporal description of the ion cloud. This description can be correlated with optical data that were taken during the short optical window. The spatial and temporal description of the ion cloud will assist the interpretation and analysis of data taken in the course of the main experiment, which was the evaluation of communication links during disturbed conditions in the propagating medium.

This report describes the experimental results and a review of the data on a test-by-test basis. The calibration ritual that took place before each experiment will be described in Section II. Sections III, IV, V, and VI will be devoted to the individual tests.

II PRELIMINARY MEASUREMENTS

At the beginning of each test and about 30 min before the launching of the barium payload, a series of measurements designed to calibrate the radar were made. There were two kinds of calibration measurements made: system noise measurements and ionospheric measurements.

Histograms of the system noise and of the system noise plus noise source were made by sampling the ADC output of the radar. Approximately 10^6 samples were obtained for each polarization channel. The ADC output was read directly into the computer through the radar-computer interface (RICE). The numbers obtained from the ADC are those that were translated into electron densities during the cloud tracking exercise.

Figure 1 is an example of the histograms of the system noise and the resulting distribution functions of the horizontal and vertical system noise outputs. These histograms were repeated for every day of the PLACES series, with the same results indicating good consistency in the performance of the equipment, the horizontal channel histogram agrees very well with the theoretical curve. The theoretical curve results from assuming that the radar front end noise is a Gaussian-distributed amplitude and an ideal logarithmic amplifier, and it is given by the following formula:

$$F(z) = \ln(10) 10^x e^{-10^x}$$

$$x = z - A$$

$$A = \text{Log}(P_0)$$

The vertical channel histogram, on the other hand, shows anomalous spikes at certain equally spaced levels. This is definitely produced by a malfunction in the ADC that persisted throughout the PLACES series. Fortunately, the effect of this malfunction was not very serious and was compensated for by the calibration and the averaging during the tests.

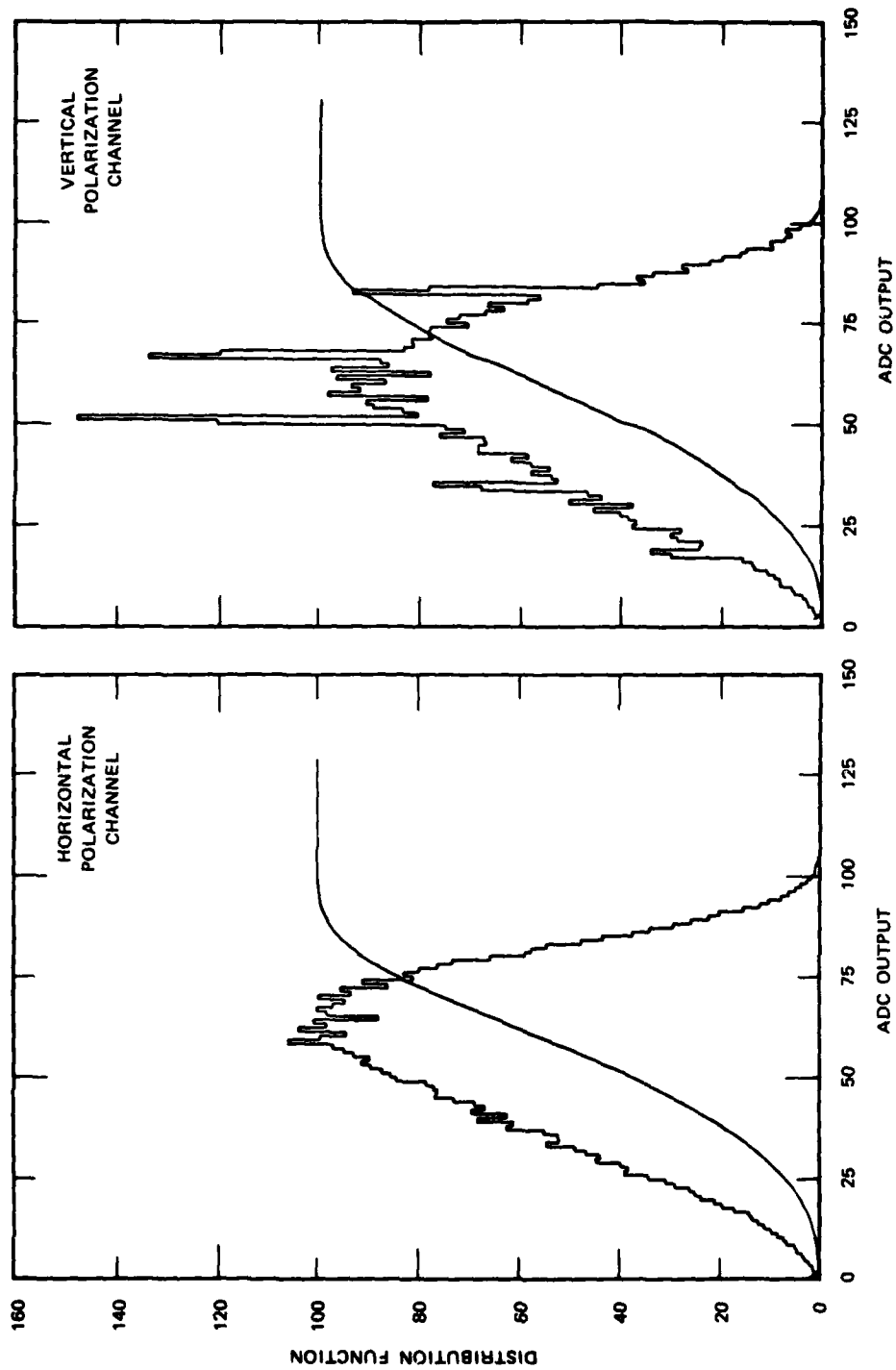


FIGURE 1 SYSTEM NOISE HISTOGRAMS AND DISTRIBUTION FUNCTIONS — EVENT GAIL

The calibration procedure itself is shown in Figure 2. A correspondence is formed between the ADC output numbers and a scale of power normalized to the average front end noise. The basis for the correspondence is equal distribution function values for the measured curved (of Figure 1) and for the theoretical power distribution curve given by:

$$F(p) = 1 - \exp(-p/P_0) \quad .$$

To clarify the procedure, let us assume that N_1 is the ADC number corresponding to a measured distribution function level, F_1 . Solving the formula above, we determine the normalized power level p_1/P_0 :

$$p_1/P_0 = -\ln(1 - F_1) \quad .$$

In this manner the calibration curve (c) of Figure 2 is formed. Actually, the computer computes and stores the power levels for each ADC level as a look-up calibration table.

This type of calibration is advantageous because it can calibrate output readings below the average noise level. In an incoherent-scatter-type measurement, this calibration is necessary because a large percentage of the samples result in measurements below the average noise power level.

The second type of measurement taken before each test is of the ambient ionosphere. Ionosphere profiles, such as shown in Figure 3, are taken by integration returns from the ionosphere for periods of 2 to 10 min according to the available time. A comparison with the maximum value of f_oF_2 values obtained with an ionosonde provides a point of reference in the conversion of normalized power to absolute electron densities. Ionosondes are able to measure the critical frequency of the ionosphere with a high degree of accuracy so that the measured constant relating the returned power to the electron density is very reliable.

Figure 4 summarizes the f_oF_2 values measured during the hour preceding each event, and verbally transmitted over the experimenter's network.

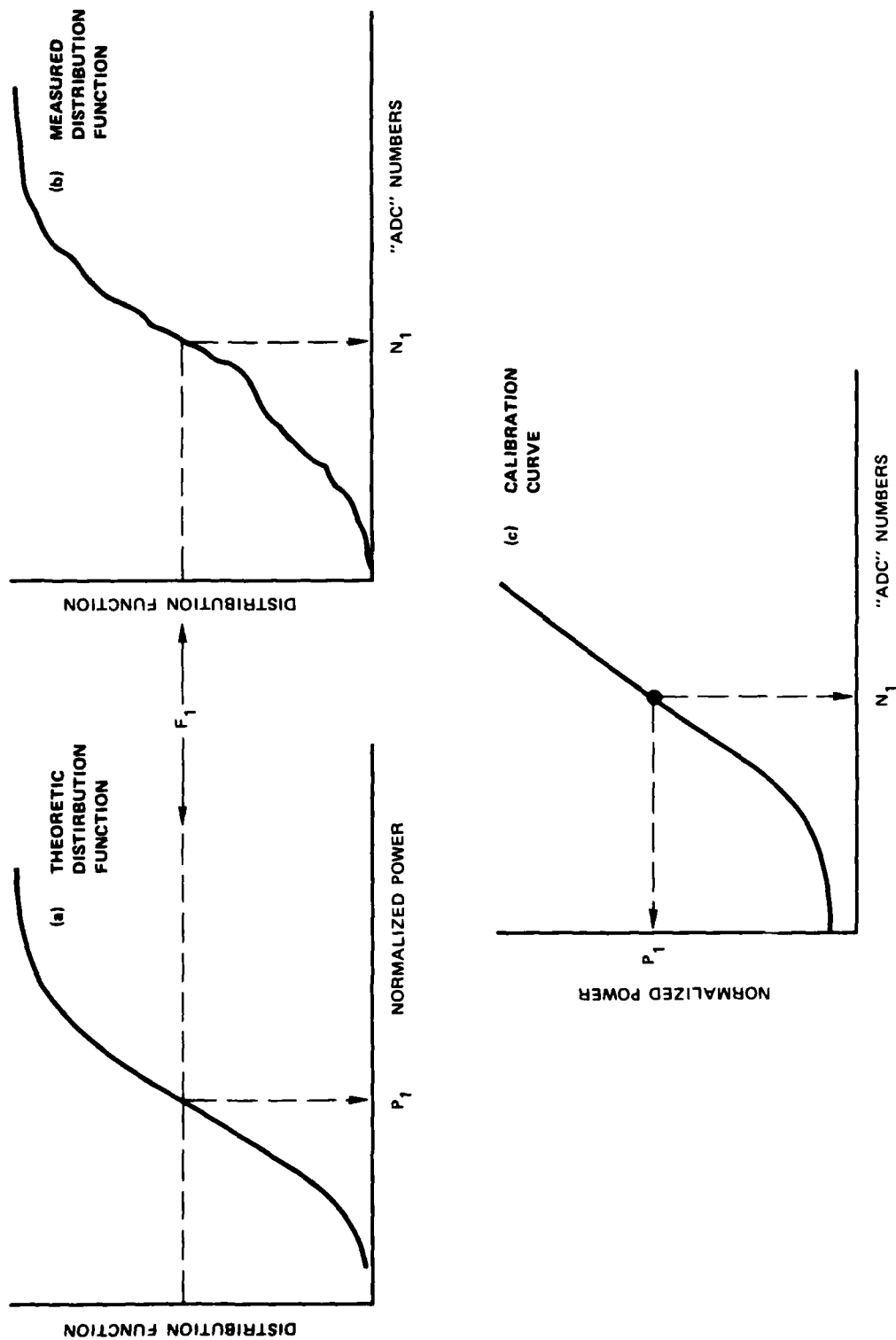


FIGURE 2 USE OF THE NOISE HISTOGRAM TO CALCULATE THE RECEIVER CALIBRATION CURVE

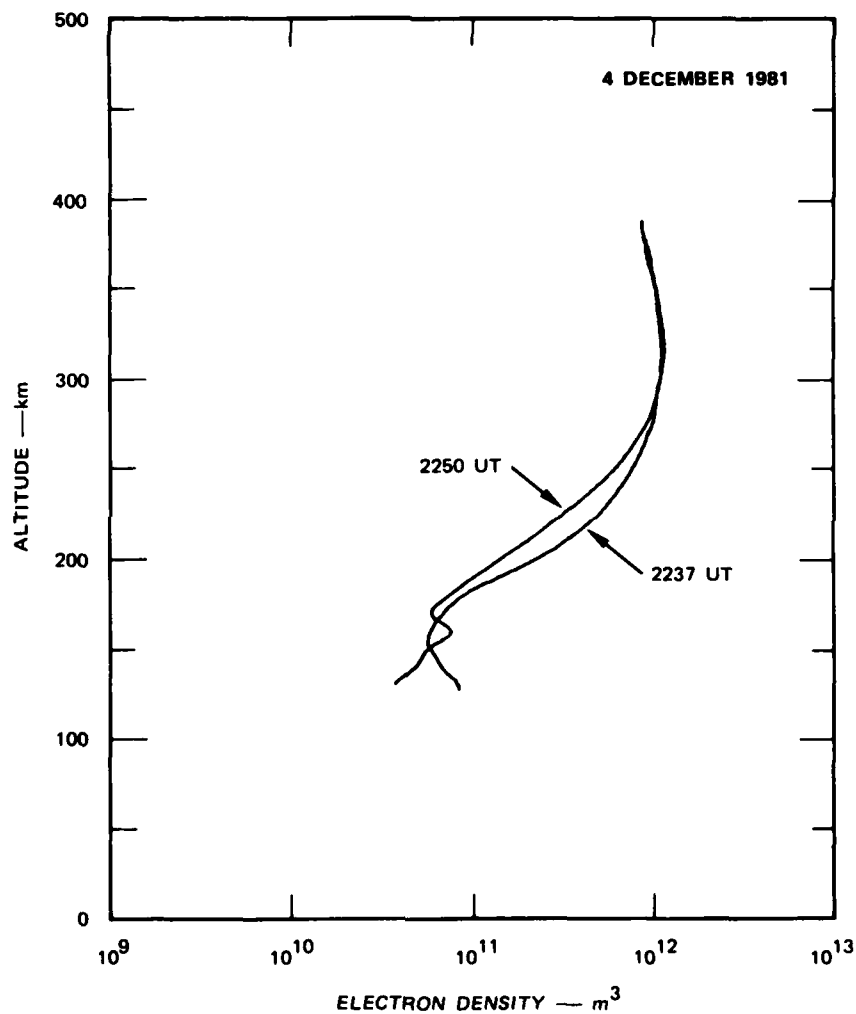


FIGURE 3 IONOSPHERIC PROFILES OBTAINED BEFORE LAUNCHING THE PAYLOAD FOR EVENT GAIL

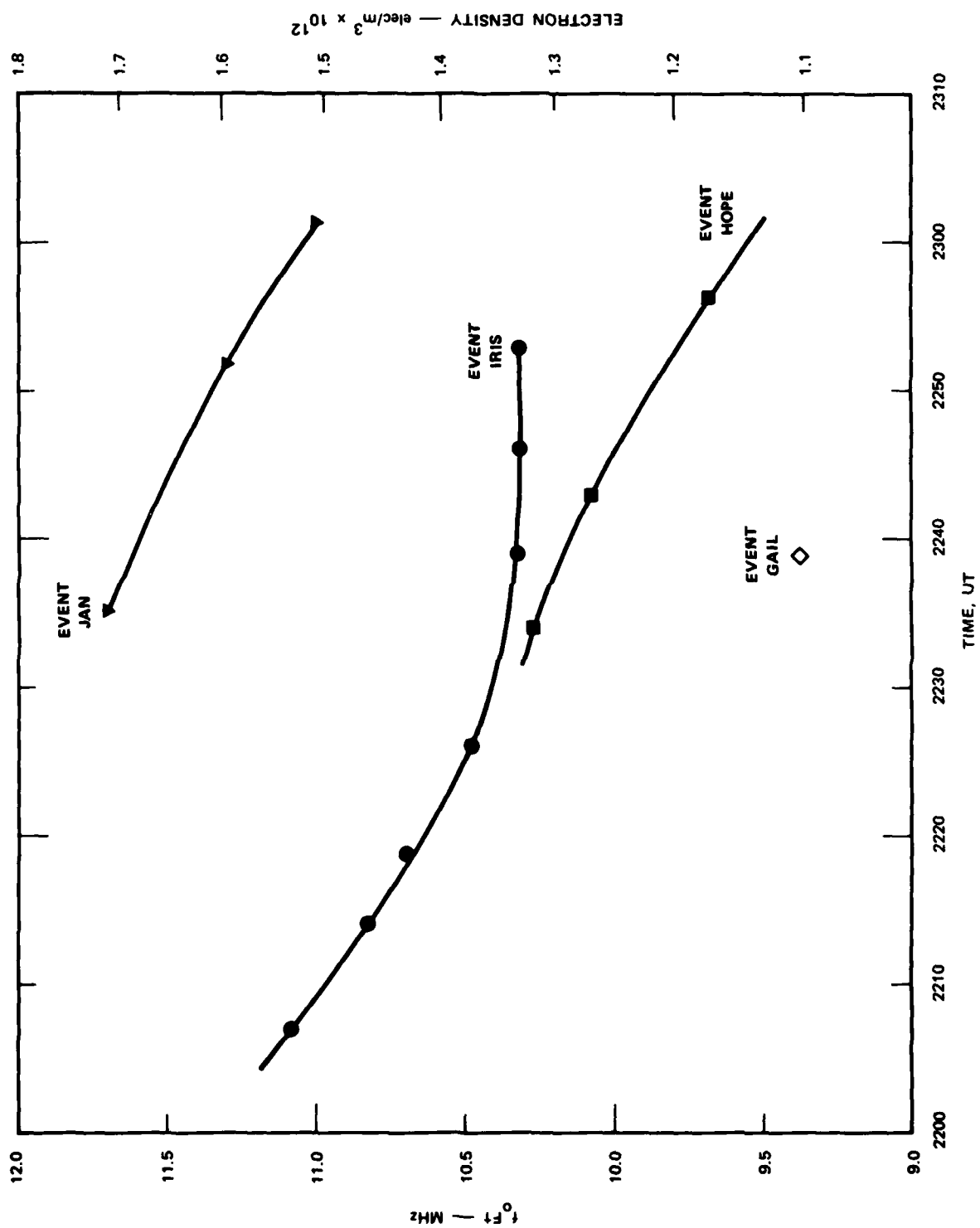


FIGURE 4 f_oF_2 VALUES READ FROM THE IONOSONDE IN REAL TIME

III EVENT GAIL

Event GAIL was tracked for approximately two hours after release. Figure 5 shows the small motion followed by the tracked point of the ion cloud. Except for the first few minutes after release when the track was still not well established. The drift was eastward up to about $T + 45$ min and westward until the track was interrupted. The northbound motion on the cloud could be caused by a descent of the ion cloud along the magnetic field line or it may be caused by the change of the position of the point with the largest electron density within the cloud itself.

The change of altitude of the ion cloud is shown in Figure 6. The cloud has dropped rapidly, and in two hours descended to an altitude of 125 km. The average vertical rate of descent was 7.8 m/s, and the initial downward velocity was 23 m/s.

Figure 6 also shows, by the spread of the data points in altitude, the times at which the tracking of the ion cloud was good and the times at which it was not. From $T + 20$ min until $T + 130$ min, the vertical spread of altitude data was very narrow, indicating that a good track was established during this period. Before the $T + 20$ min mark, a very wide spread in the data points indicates problems in the tracking process. A cable in the radar controlling the sampling of data broke at about the time of release. Data and track were lost during the time necessary to diagnose and solve the problem. When the radar finally began following the proper sequence of measurements to track the cloud, it took some time before zeroing in on the point of greatest electron density.

The time history of the maximum electron density is shown in Figure 7. The small spread of data after $T + 20$ min is also evident on this graph. The spread in electron density measurements is caused by random errors and by pointing errors. The statistical errors occurred because a limited number of pulses was used for integration at each antenna beam position and these statistical errors were the limit of the accuracy for the

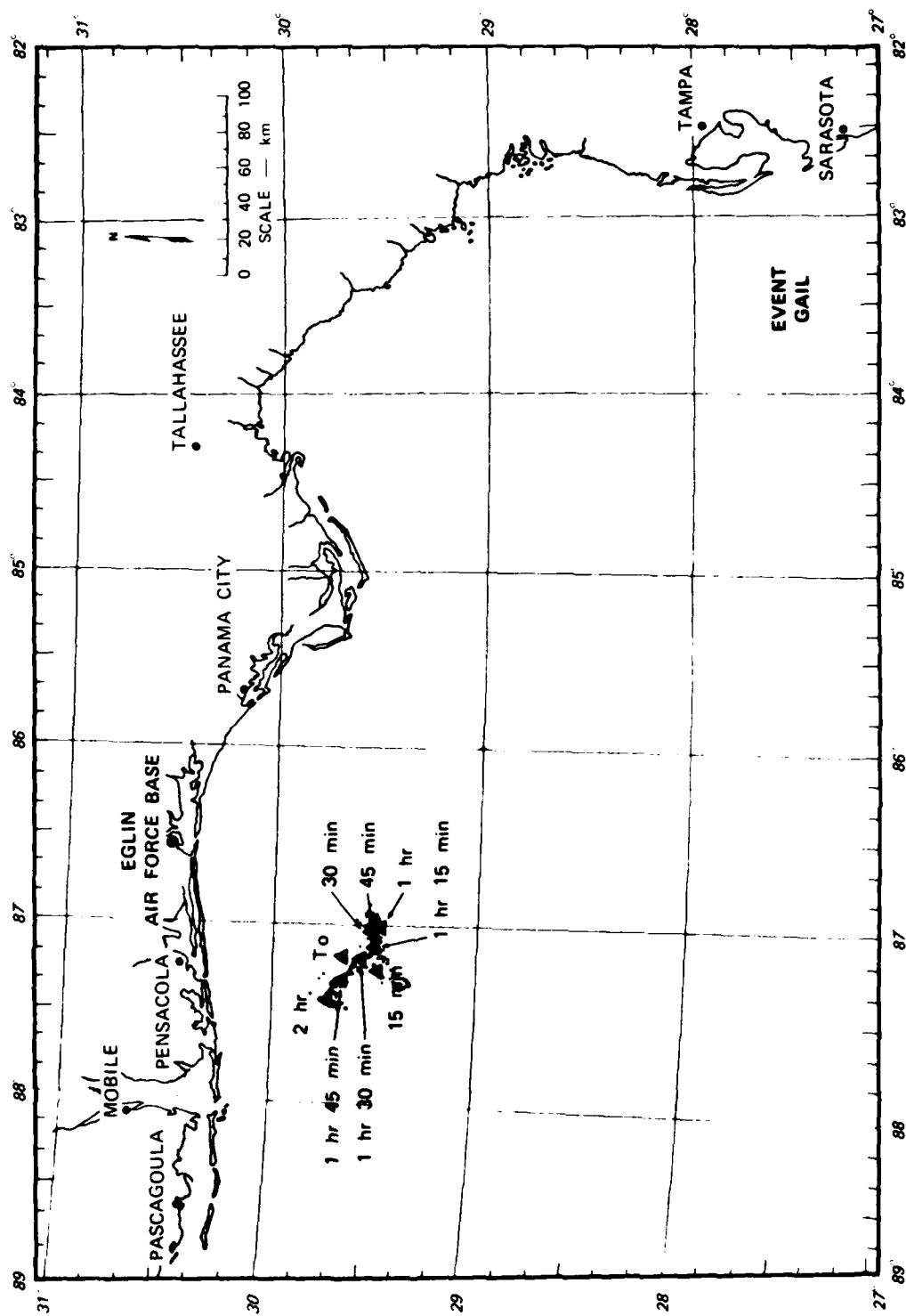


FIGURE 5 HORIZONTAL TRACK OF EVENT GAIL

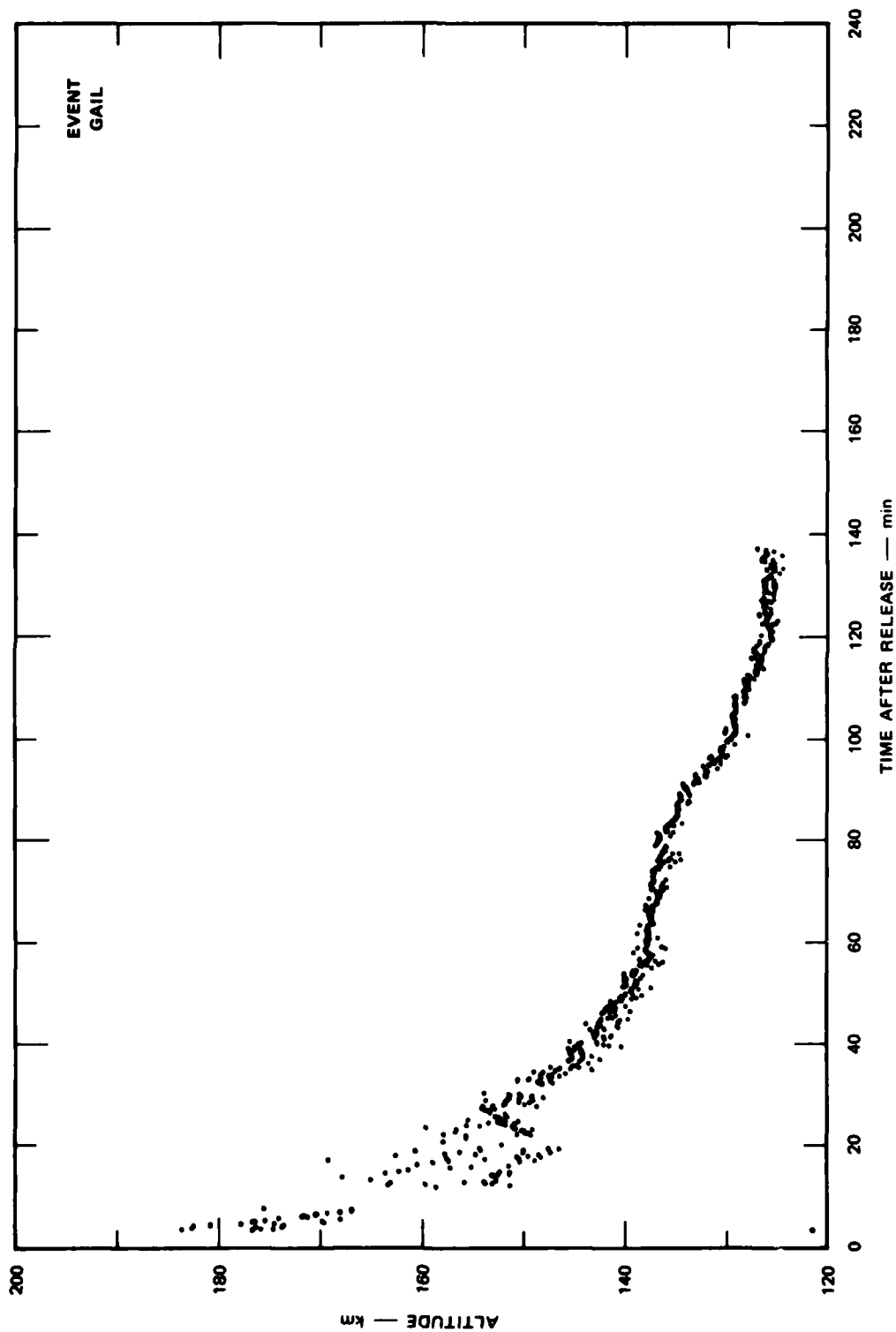


FIGURE 6 ATTITUDE DATA FROM EVENT GAIL

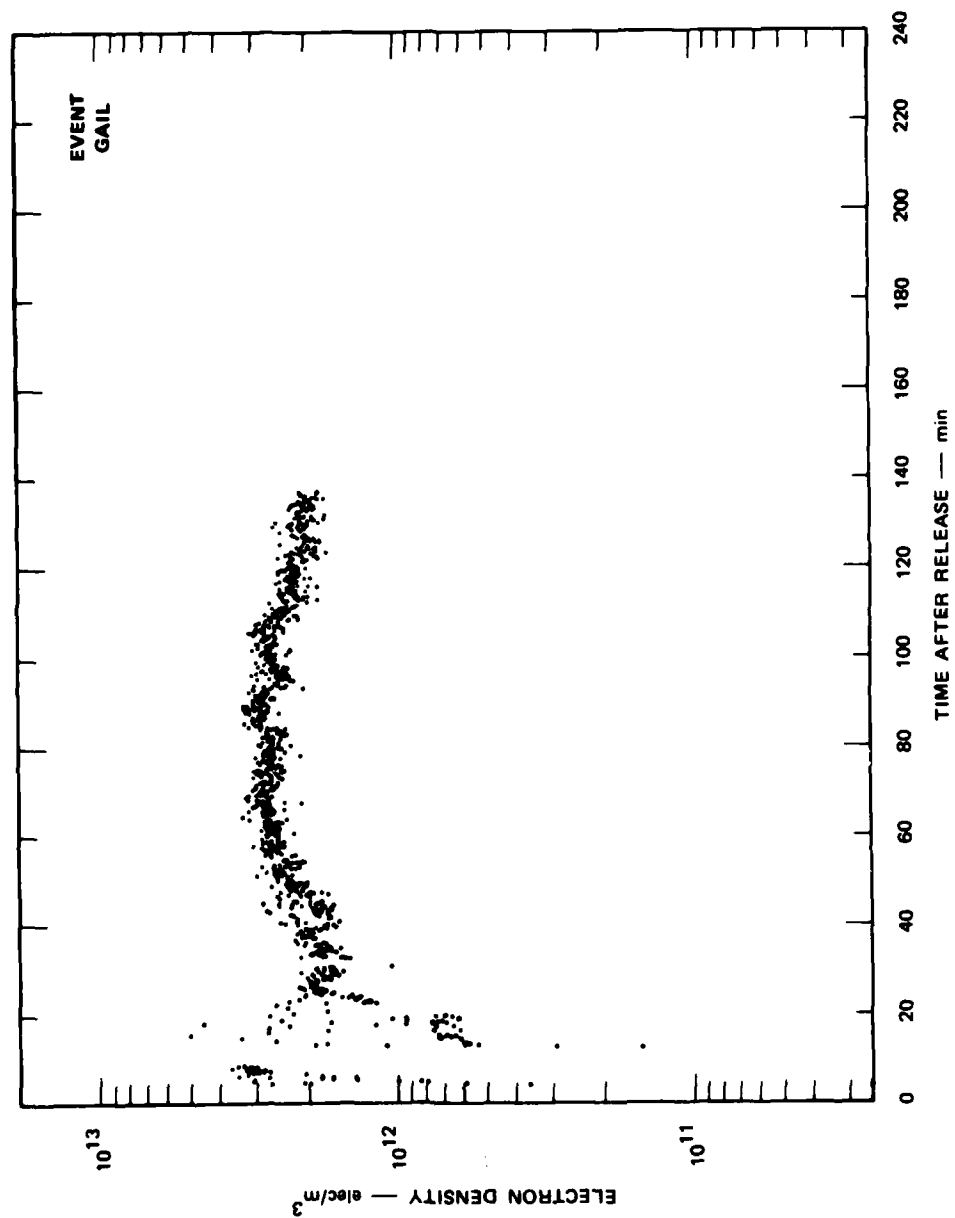


FIGURE 7 MAXIMUM ELECTRON DENSITY AS A FUNCTION OF TIME FOR EVENT GAIL

electron density measurements. The pointing errors on the other hand reflect a limitation in the tracking algorithm which did not always point the antenna beam to the point of maximum electron density. The pointing errors depend on a complicated set of geometric considerations and is difficult to describe; however, they appear to be larger than the random errors in some circumstances. The total spread of points in Figure 7 has a sigma of about 4 to 5%. The main limitation in the electron density measurement is the accuracy of the absolute calibration of the radar.

IV EVENT HOPE

Event HOPE was the smoothest running test from the point of view of radar tracking. The ion cloud was tracked for over 4 hours after release, and good quality data were obtained. The ground track of Event HOPE is shown in Figure 8. There is a monotonic westward motion throughout the life of the ion cloud. There are some reversals of motion in the North-South direction, however, that should be reviewed with care. The southward motion is unquestionably a drift involving ionization that crosses magnetic field lines. The motion to the north may be caused by the vertical fall of the ion cloud with a tendency to slide northward along the magnetic field lines. As has happened in other experiments, the radar sometimes changes the tracked point from one region of the cloud to another, and this tendency could be responsible for the apparent reversal between $T + 60$ min and $T + 75$ min.

The vertical descent of the ion cloud shown in Figure 9 indicates a rate of descent of 11 m/s during the first 50 min. The vertical velocity then slowed to only 2.5 m/s and the ion cloud dropped to an altitude of 120 km at four hours after release. Understanding the evolution of an ion cloud and following its changes when it reaches such a low altitude is interesting to researchers on this field.

The maximum electron density shown in Figure 10 shows a remarkably constant value after the first 20 min. This near constant value occurs as the ion cloud changes its altitude by several kilometers. To have confidence in the altitude history of the ion cloud, we have closely examined the quality of the data acquired at times separated by three hours. Figure 11 shows the received power data acquired with 2-s or 80-pulse integration and plotted as a function of altitude. We see that even though the returns from the ion cloud at 227:27 UT are five times smaller than at 2311:51 UT, the returns from the ion cloud are still clearly well defined and definitely above the noise level. Figure 12

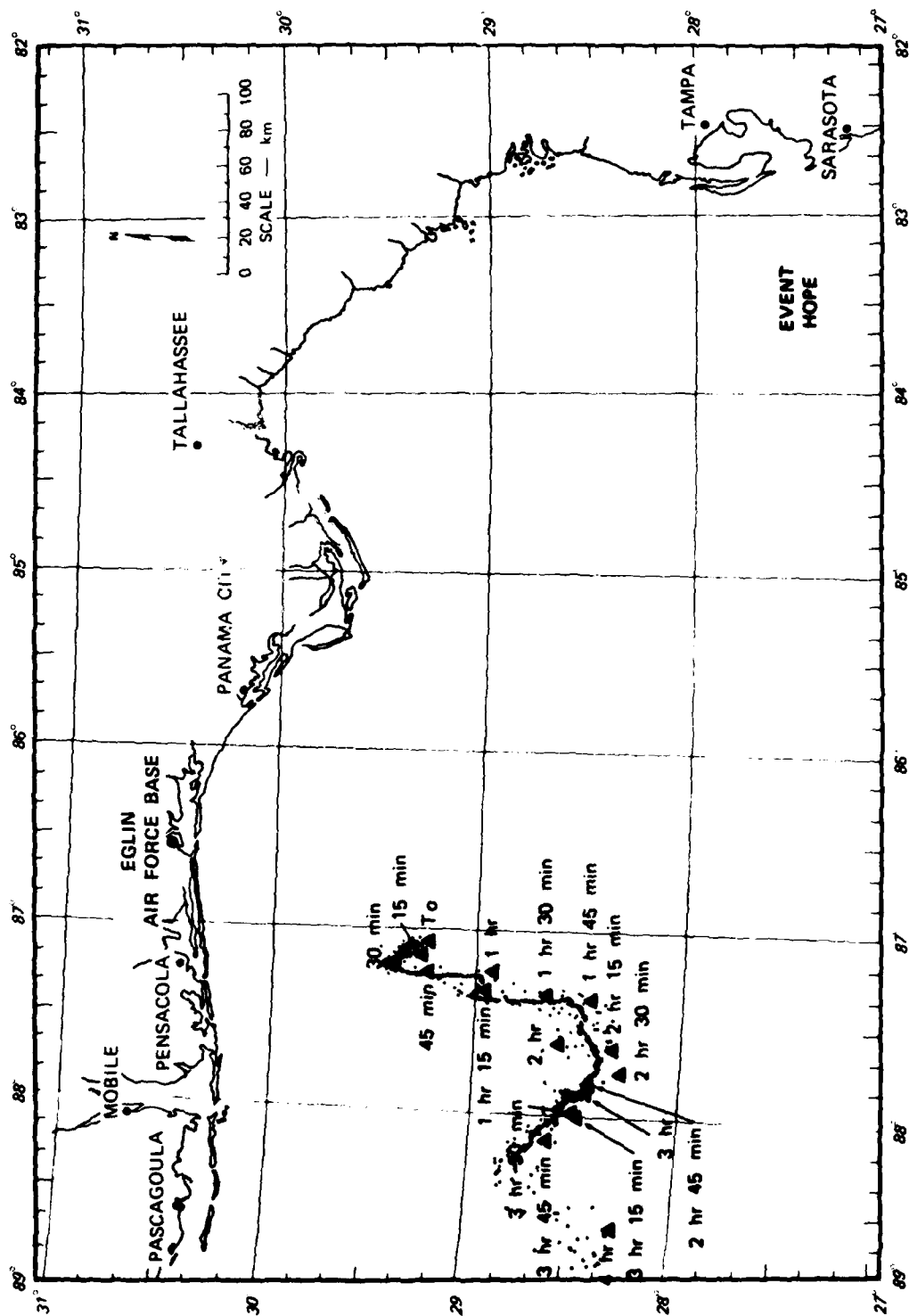


FIGURE 8 HORIZONTAL TRACK OF EVENT HOPE

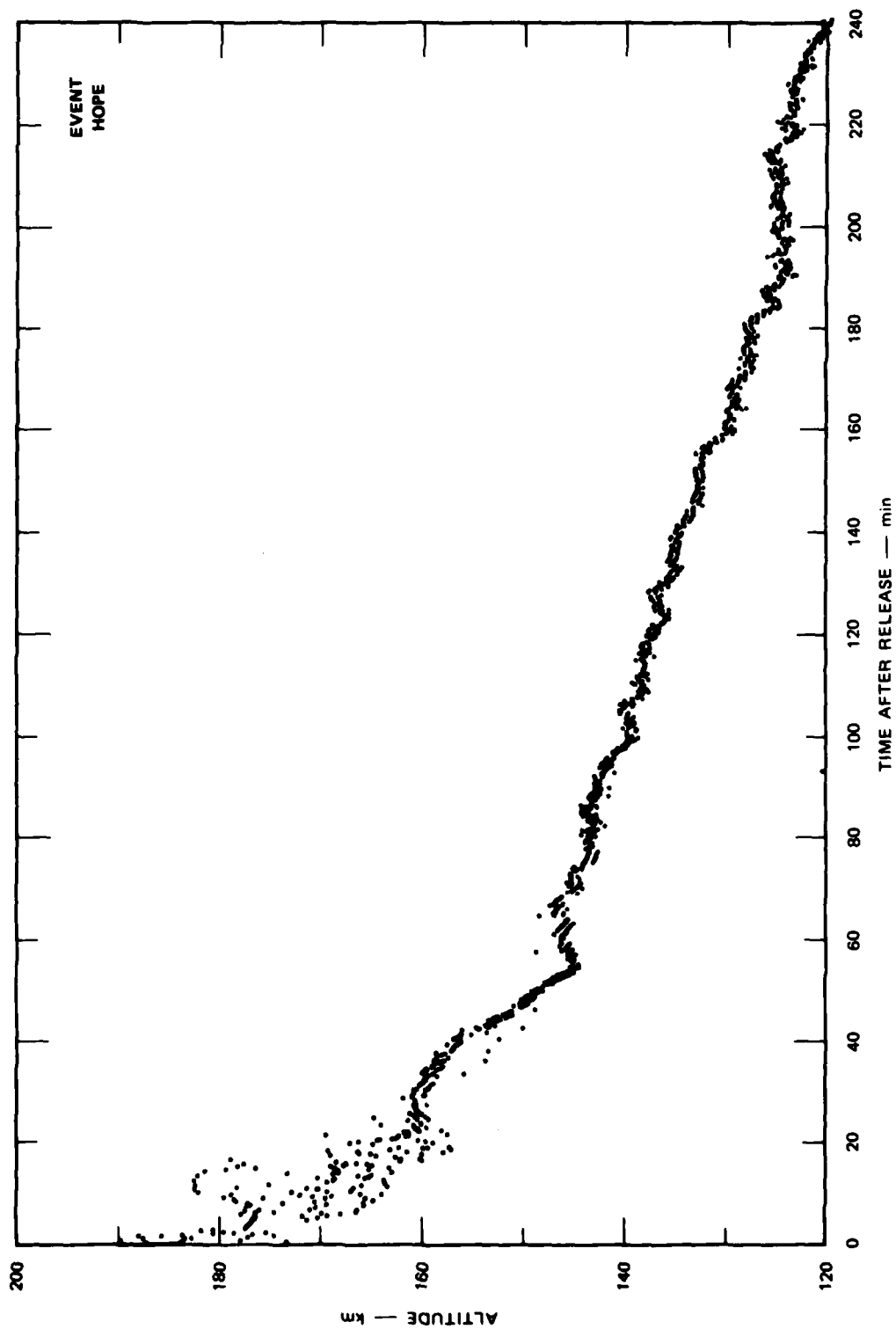


FIGURE 9 ALTITUDE DATA AS A FUNCTION OF TIME OF EVENT HOPE

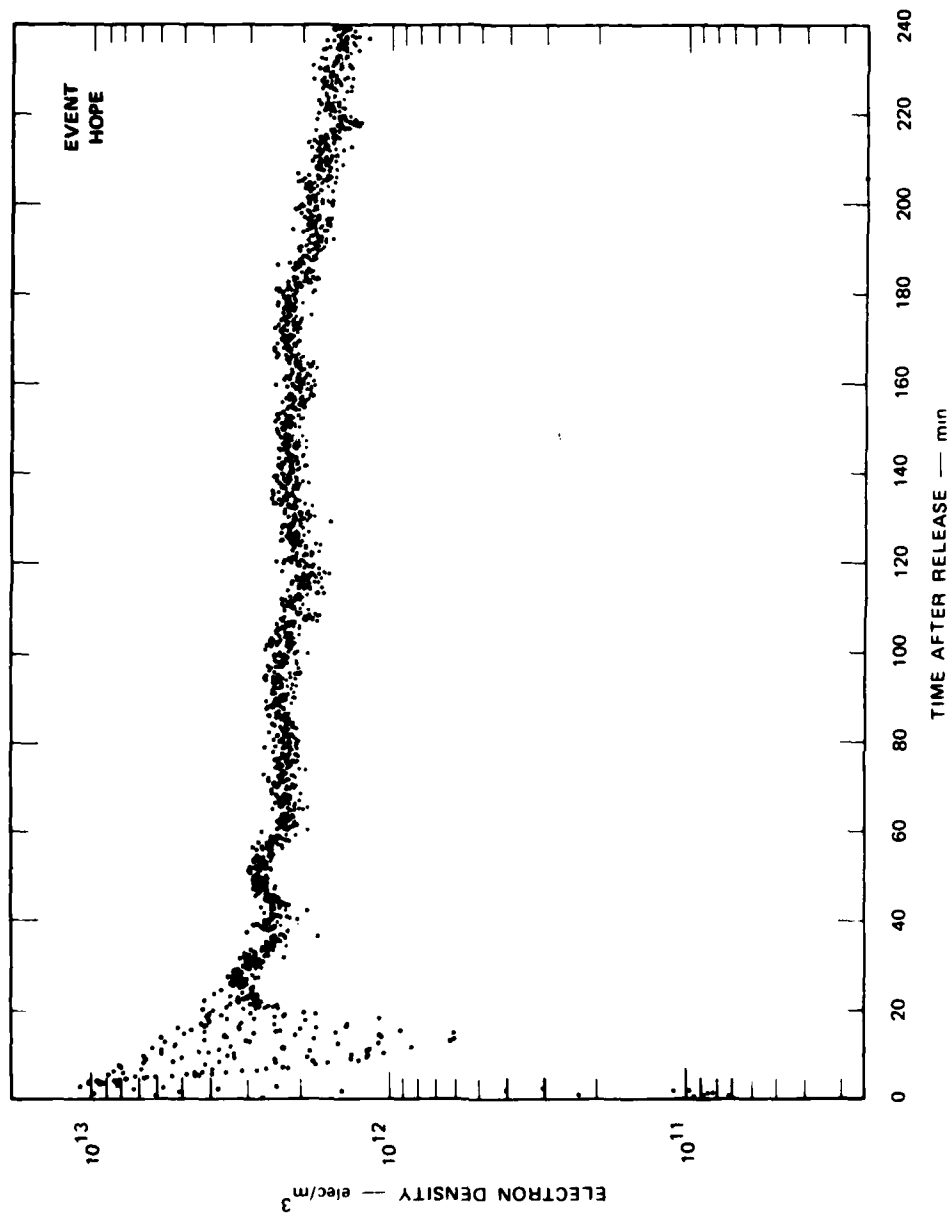


FIGURE 10 MAXIMUM ELECTRON DENSITY AS A FUNCTION OF TIME FOR EVENT HOPE

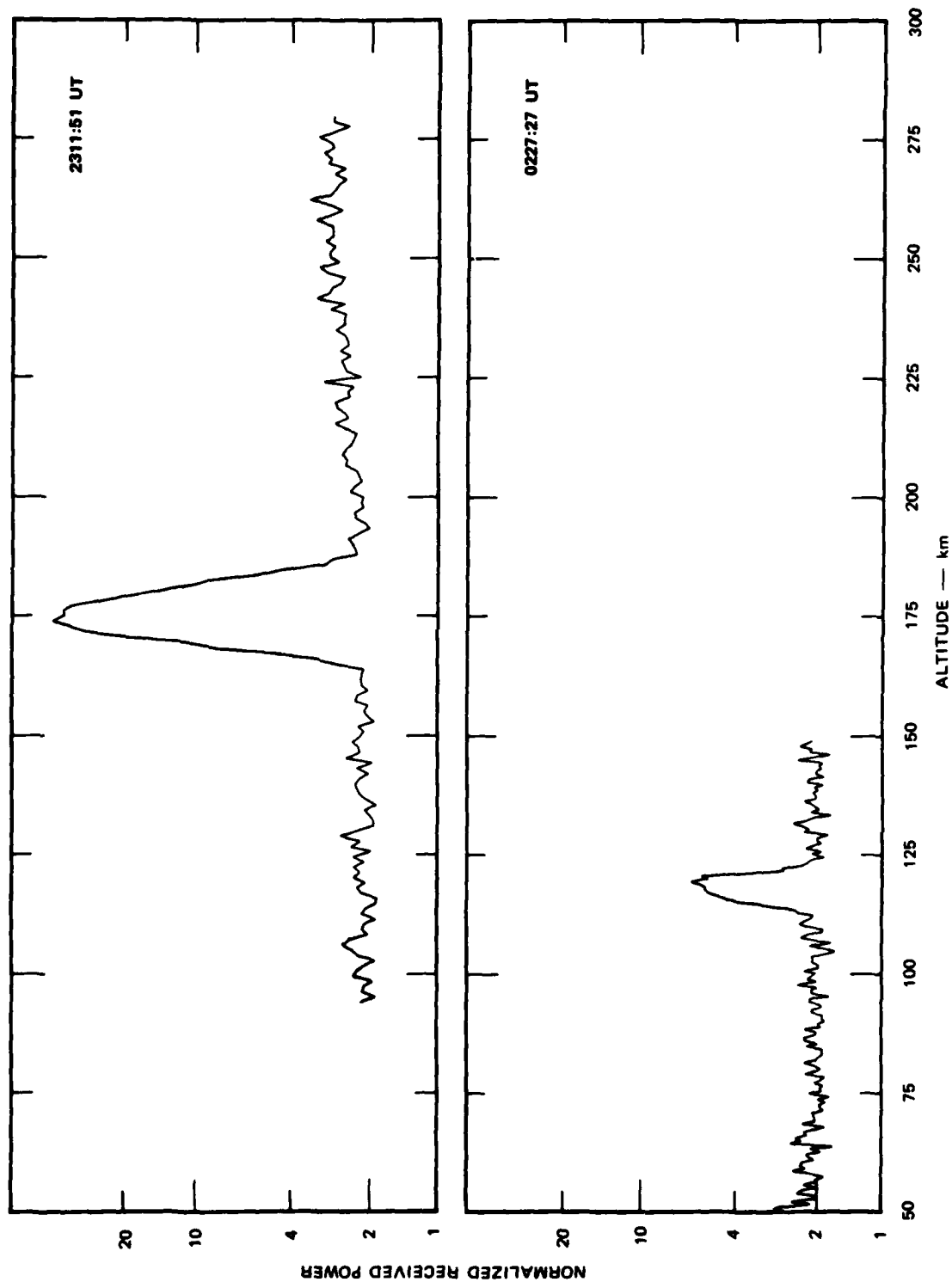


FIGURE 11 ROUGH DATA ACQUIRED FOR TWO ANTENNA-BEAM POSITIONS WITH TWO-SECOND INTEGRATION TIME OF EVENT HOPE

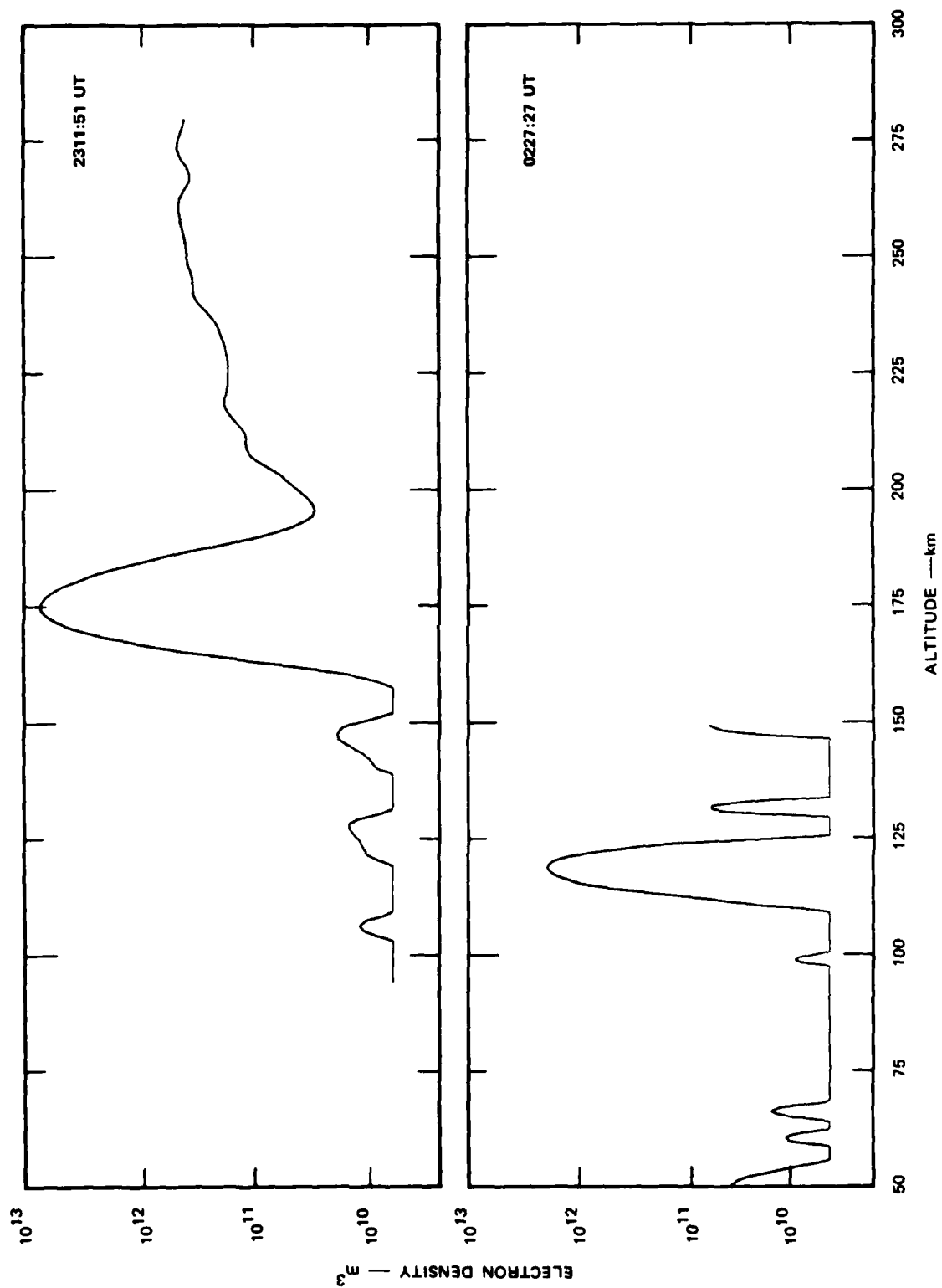


FIGURE 12 ELECTRON DENSITIES DERIVED FROM DATA IN FIGURE 11 FOR EVENT HOPE

shows the same data after the system noise, normalized to 2, has been subtracted, the range correction has been made, and the conversion to electron density has been completed. The ionosphere background above 200 km is evident. The vertical profile of the ion cloud at 0227:27 UT seems narrower than that at 2311:51 UT. This may not be the case, however, because the radar beam traversed the ion cloud at a very low-elevation angle and the apparent narrow vertical extent of the ion cloud may be caused by the radar beam entering through one side of the cloud and leaving through the other side of the ion cloud. In other words, the cloud may extend above and below 120 km and remain outside the radar beam. This problem should be studied in greater detail, in case another barium release program is contemplated.

Vertical profiles of the ion cloud were formed by choosing the maximum electron density at a given height from a set of approximately 40 to 50 antenna-beam positions. We can reasonably assume that these profiles show the electron density along the magnetic field line with the maximum electron content. Some words of caution should be added because the discrete number of antenna beam positions used during tracking do not assure that all the points along the magnetic field line of interest have been observed and measured. The resulting vertical profiles are shown in Figures 13, 14, 15, and 16 for various times.

The first vertical profile at 2325 UT is about 18 min after release and the 3-dB width is 20 km. As the ion cloud grows older and drops to lower altitudes, the three-dB width narrows to 12 km at 2404 UT, to 10 km at 2417 UT, and to about 7 km after 0220 UT. This decrease of the ionization in the vertical direction should be investigated further to verify whether it is caused by the geometric configurations of the observations since the horizontal range combined with the low altitude require every low-elevation pointing of the radar beam. A somewhat different view of this vertical extent of the ion cloud is seen in the next section, devoted to Event IRIS.

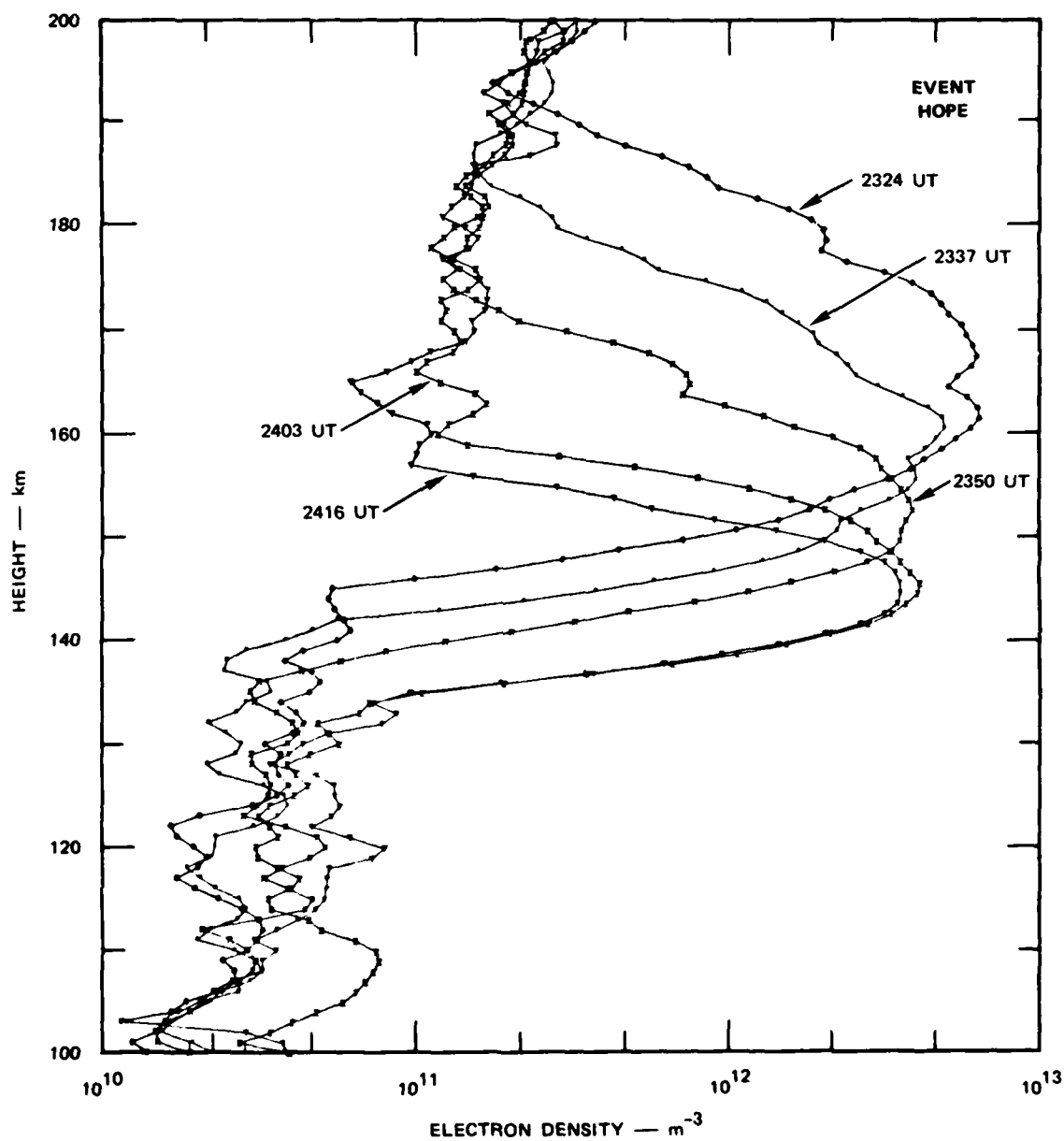


FIGURE 13 VERTICAL ELECTRON DENSITY PROFILES OF EVENT HOPE AT VARIOUS TIMES

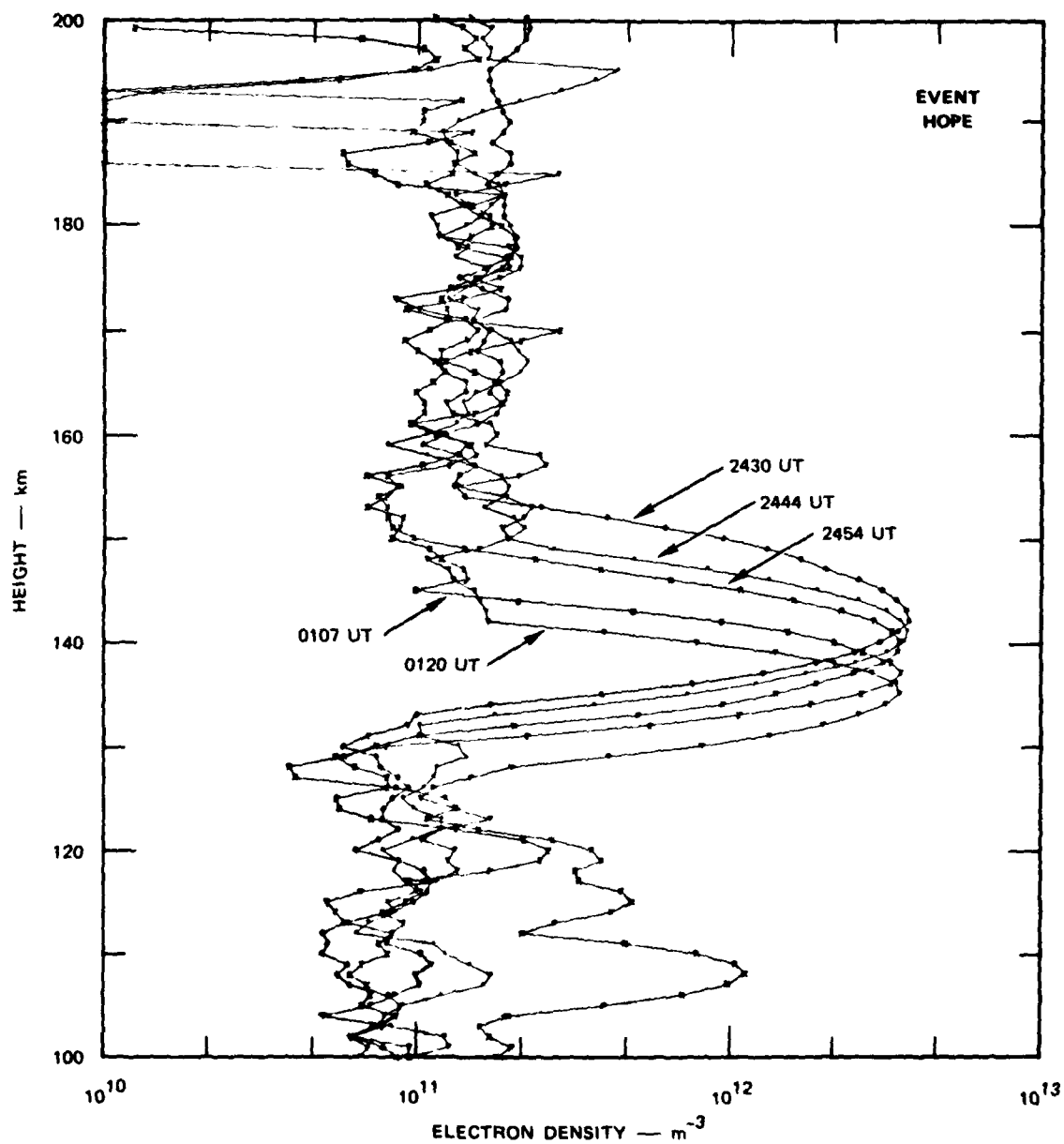


FIGURE 14 VERTICAL ELECTRON DENSITY PROFILES OF EVENT HOPE AT VARIOUS TIMES

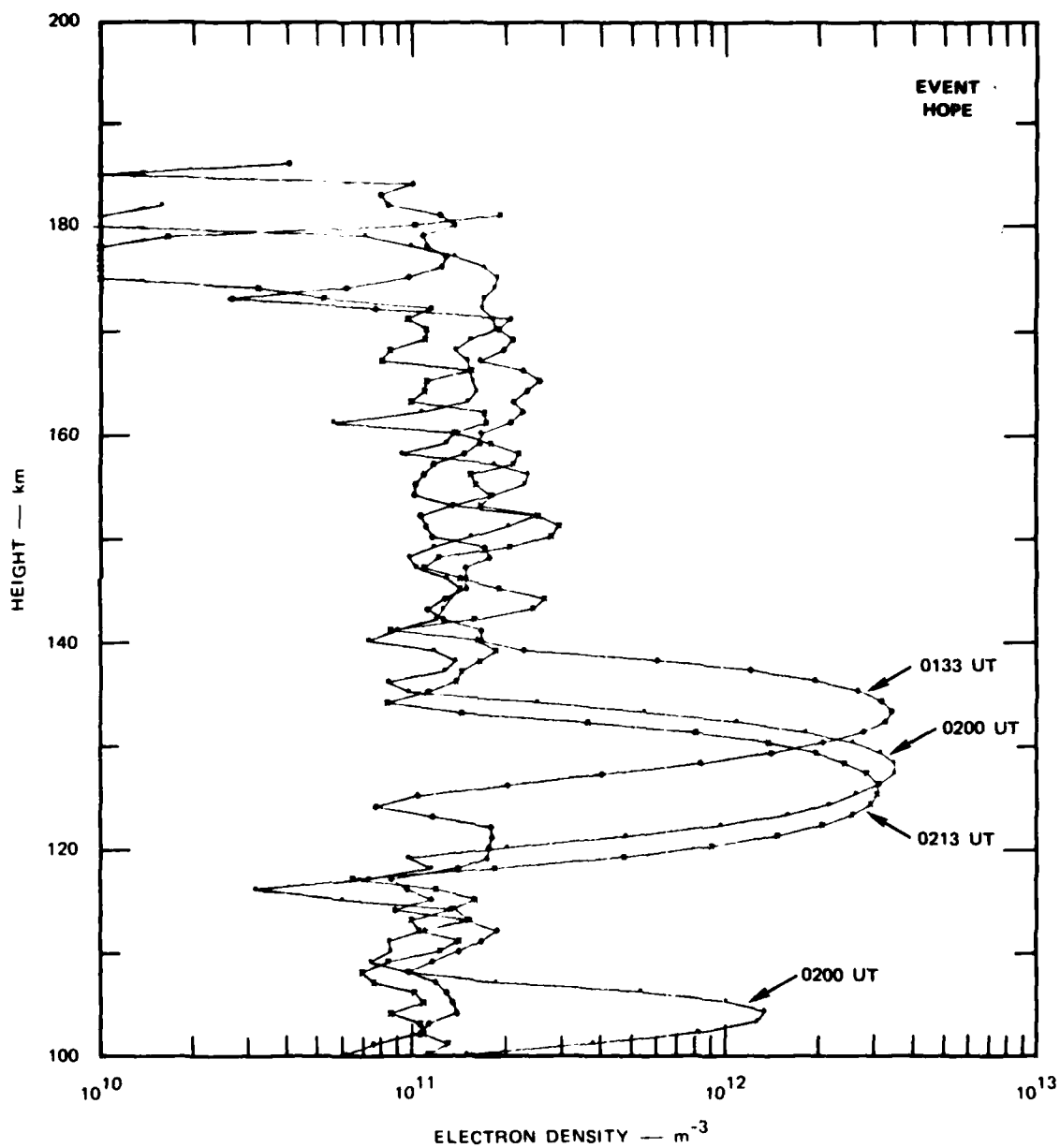


FIGURE 15 VERTICAL ELECTRON DENSITY PROFILES OF EVENT HOPE AT VARIOUS TIMES

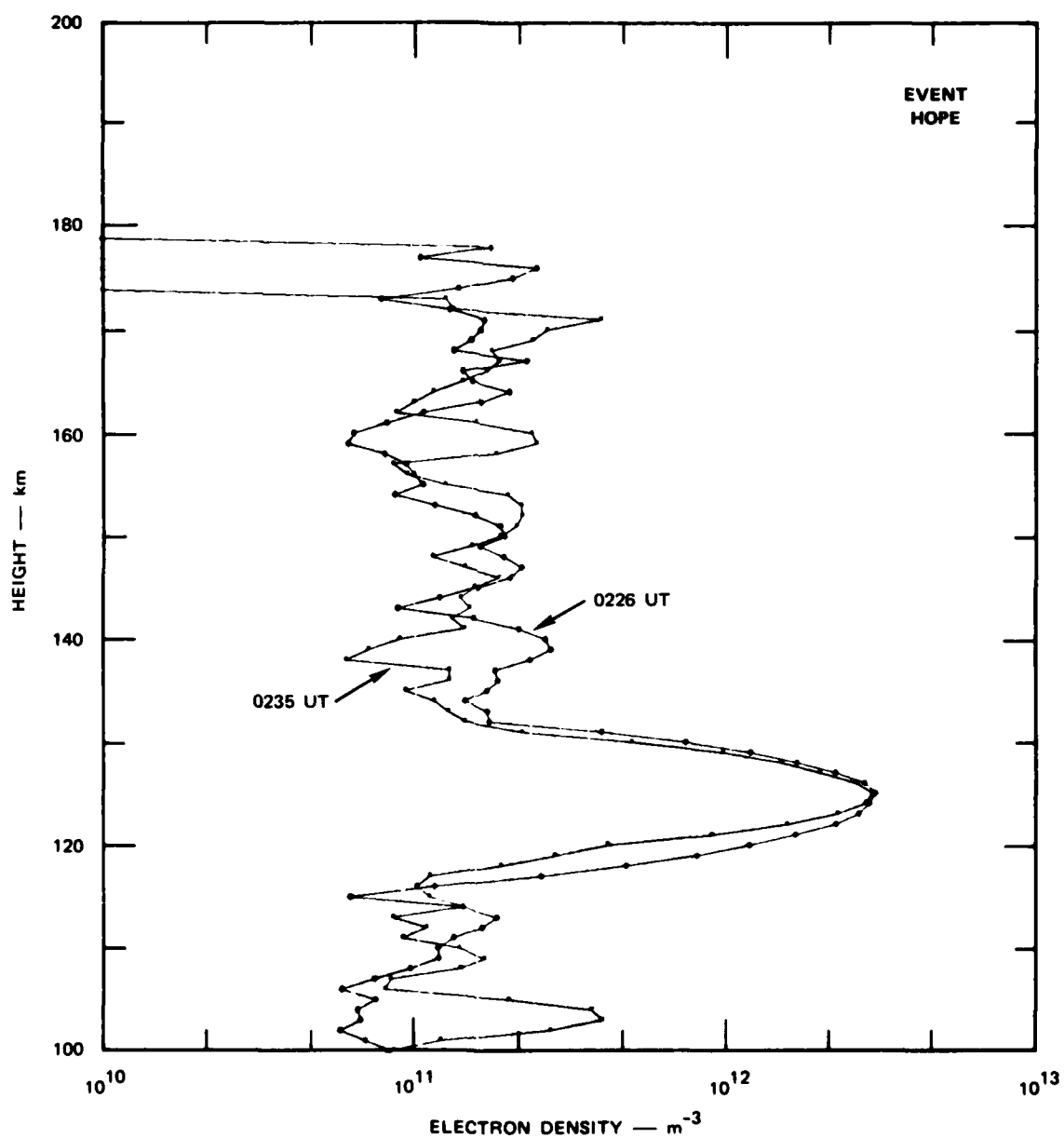


FIGURE 16 VERTICAL ELECTRON DENSITY PROFILES OF EVENT HOPE AT LATE TIMES

V EVENT IRIS

The ground track of Event IRIS is shown in Figure 17. The ion cloud location was well established about 10 or 15 min after release. The motion of the ion cloud was very steady throughout the period observation, and the average velocity was 26 m/s southward and 18 m/s eastward. This drift velocity carried the ion cloud out of the preset sampling range of the radar about 90 min after release.

The maximum electron density as a function of time is shown in Figure 18. The low electron densities during the first few minutes after release show that the radar was not pointing to the center of the ion cloud. When the ion cloud was found and the radar zeroed in on the densest part of the ion cloud, the electron density climbed up to 8×10^{12} el/m³. The drop in electron density after T + 80 min and after T + 100 min was caused by pointing problems when the cloud drifted out of range and when the operator tried to change the preset range limits for the electron density measurements. The drop in electron density between T + 20 min and T + 30 min, however, seems to be an actual change in the ion cloud itself, since the track was well established at this time.

The altitude of the ion cloud is shown in Figure 19. The anomalies in these data define the times at which problems were encountered during the tracking of the ion cloud. The data before T + 10 min and after T + 75 min are very erratic with very wide scatter in the points. The narrow spread of data between T + 10 min and T + 75 min indicates that during this period good quality data were obtained.

Vertical profiles of the ion cloud at various times are shown in Figure 20 and the narrowing of the ion cloud in the vertical direction as the ion cloud reaches lower altitudes is not as obvious for Event IRIS as it was for Event HOPE. However, the measurements of Event IRIS do not span as large a period or as large a difference in height. The vertical

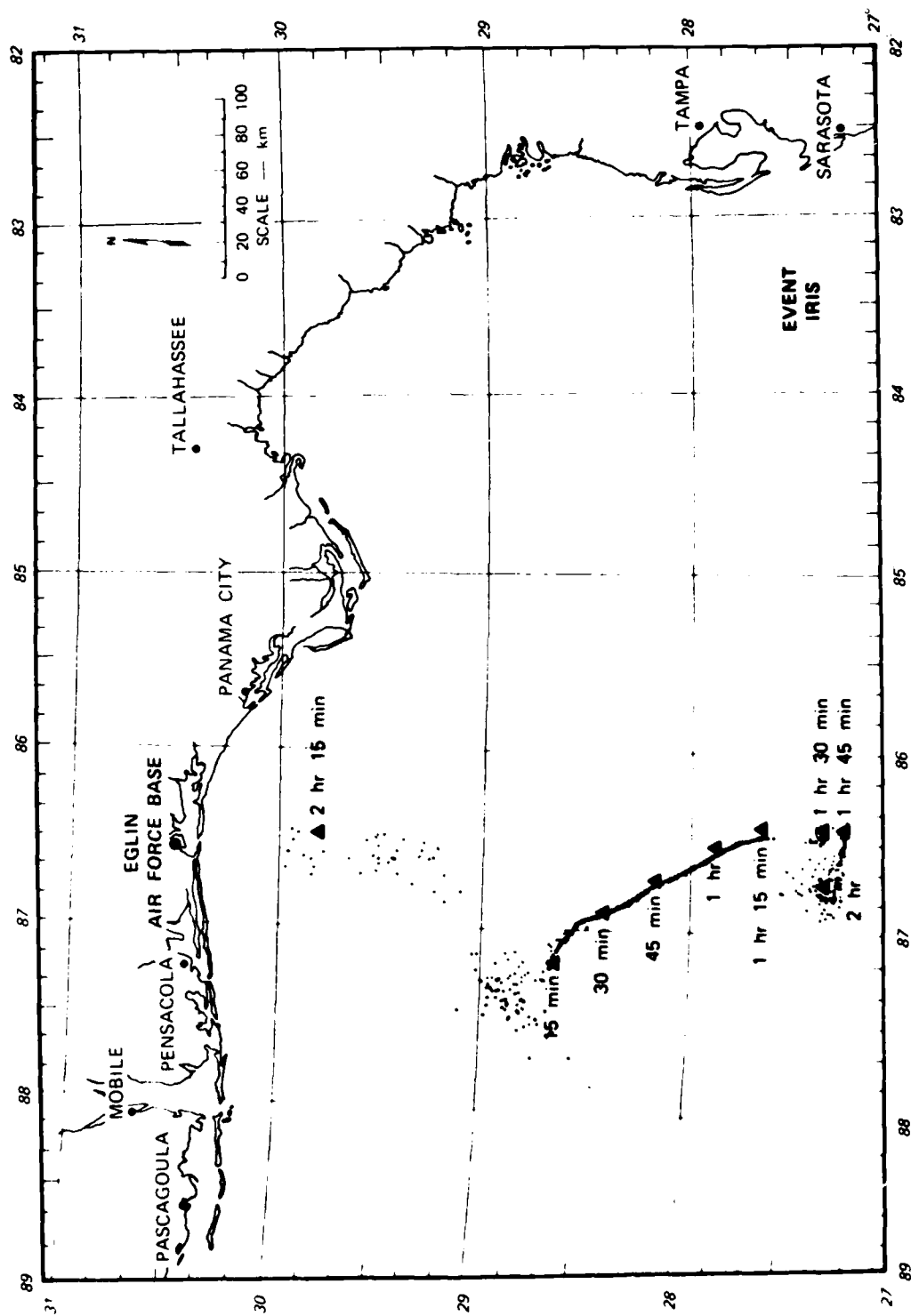


FIGURE 17 HORIZONTAL TRACK OF EVENT IRIS

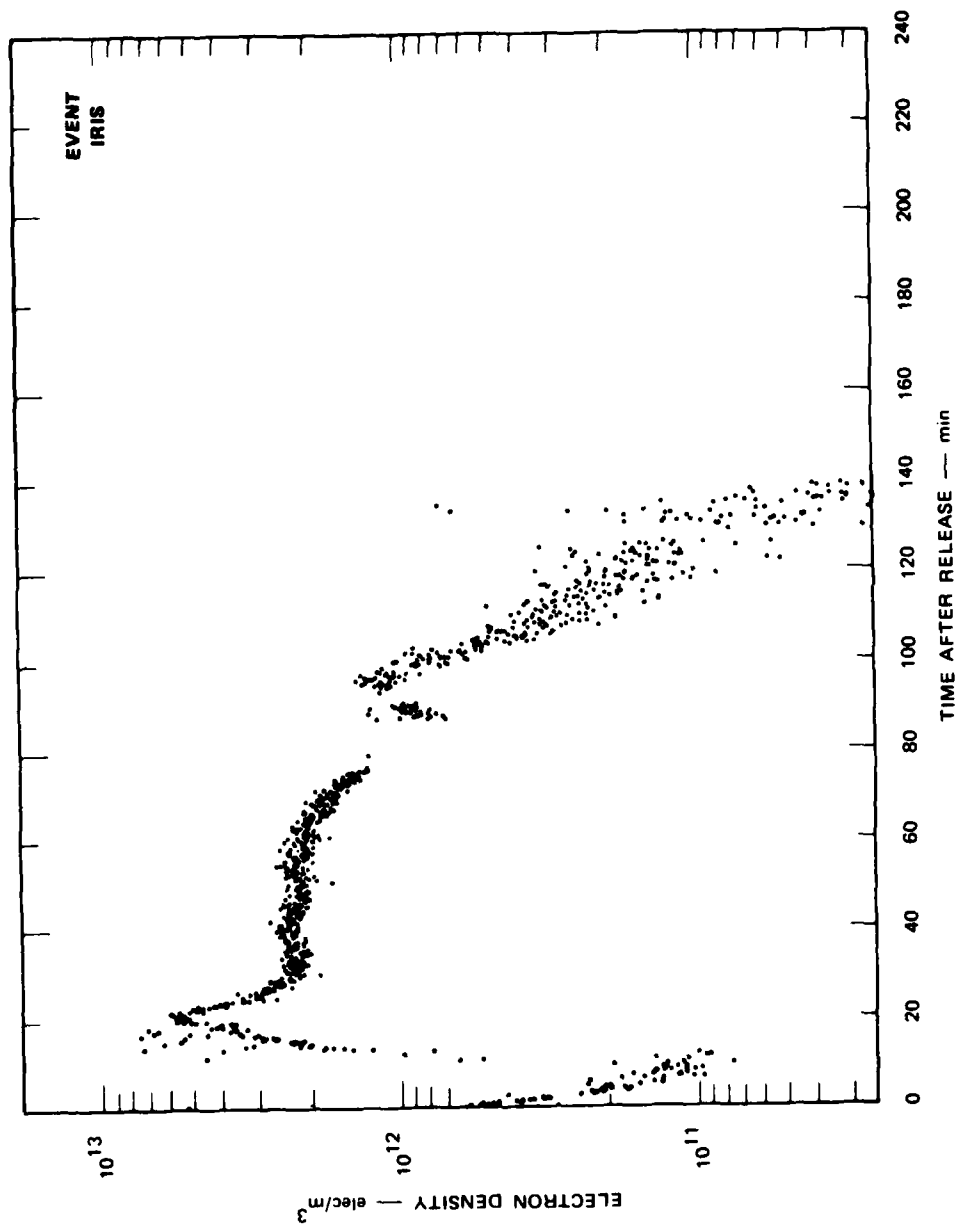


FIGURE 18 MAXIMUM MEASURED ELECTRON DENSITY OF EVENT IRIS AS A FUNCTION OF TIME

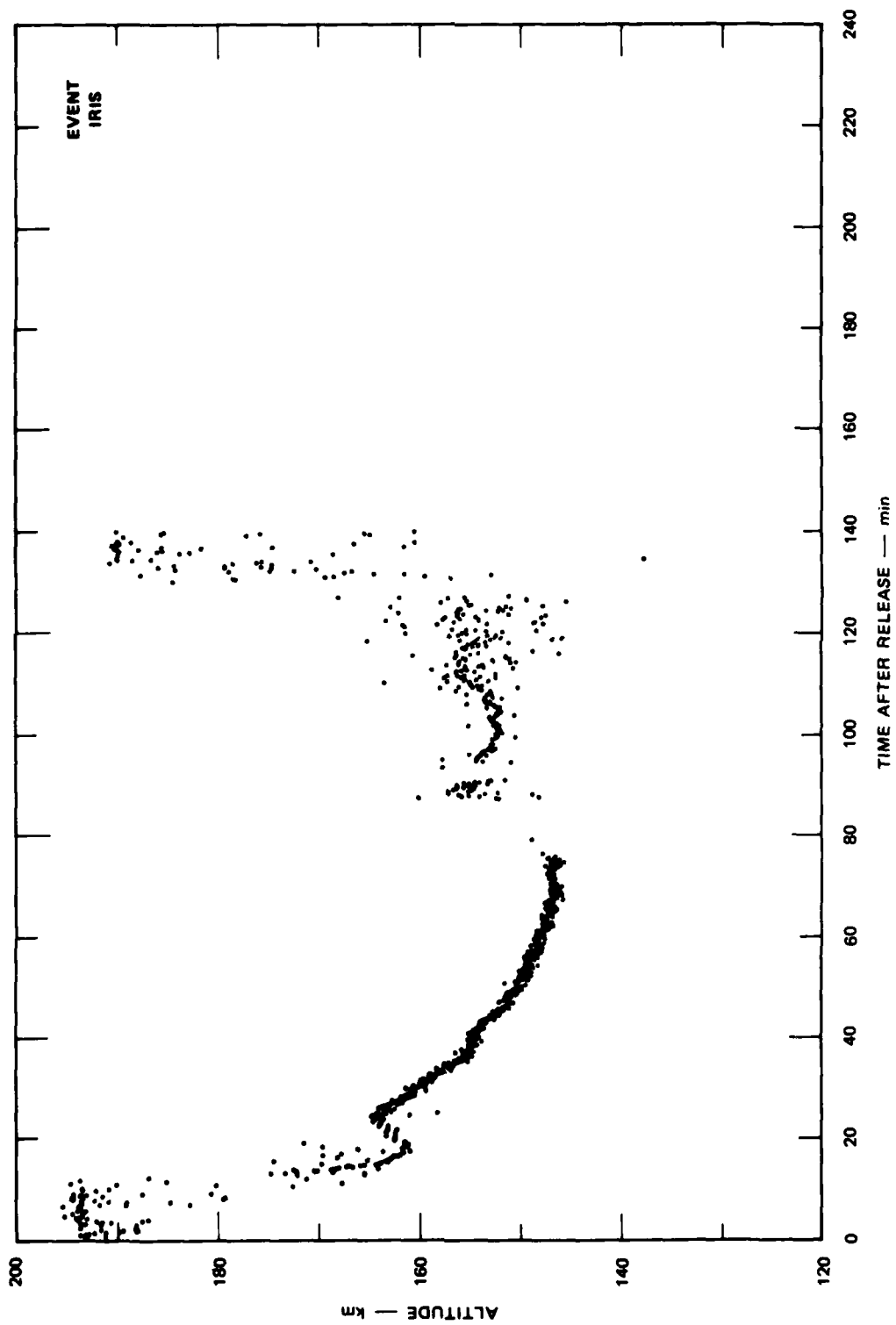


FIGURE 19 ALTITUDE DATA OF EVENT IRIS AS A FUNCTION OF TIME

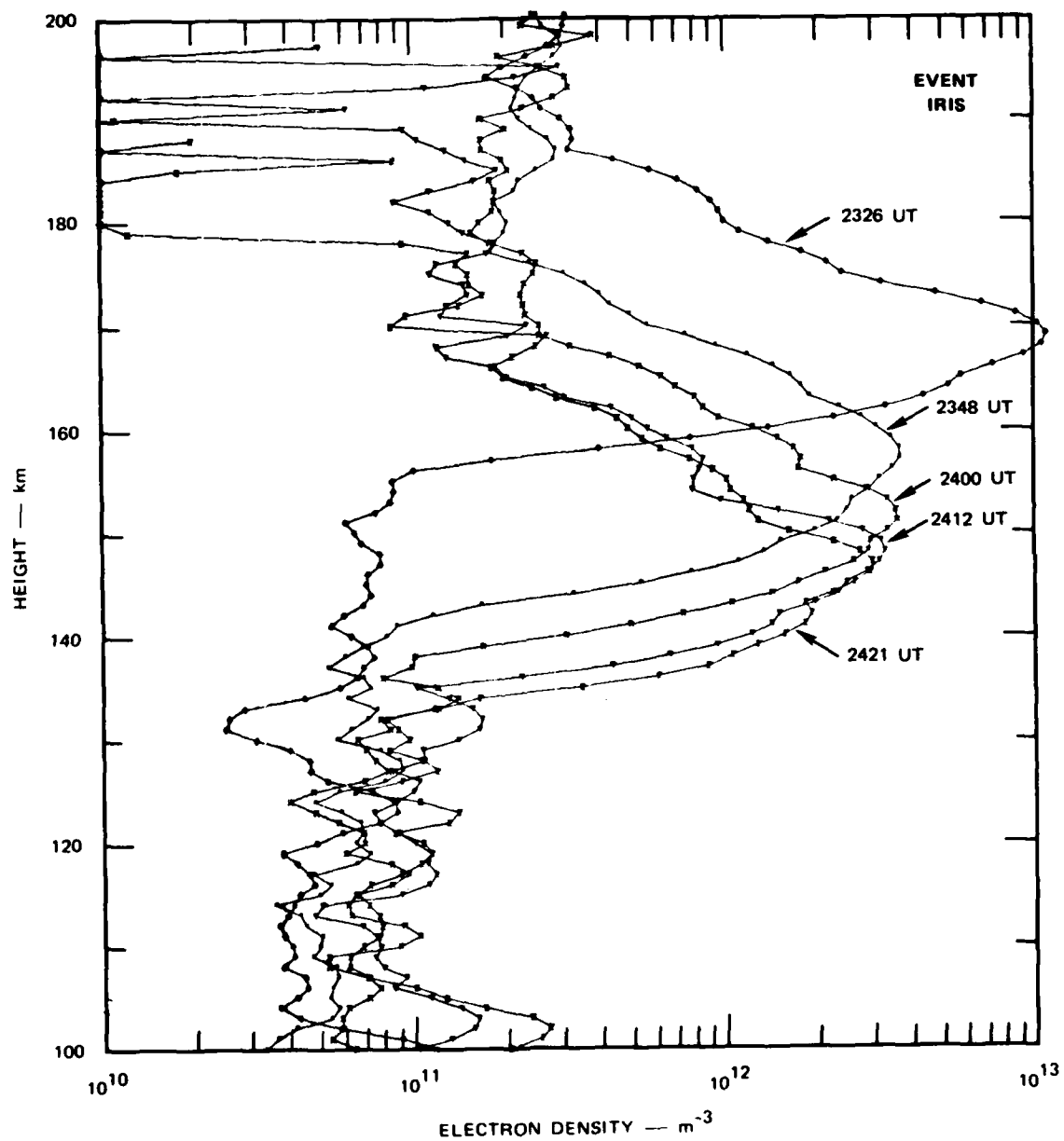


FIGURE 20 VERTICAL ELECTRON DENSITY PROFILES OF EVENT IRIS AT VARIOUS TIMES

narrowing of the ion cloud seems to be present in Figure 20, but on a smaller scale.

The horizontal shape or cut of the ion cloud has been obtained three different times, separated by about 20 min. These horizontal contours are shown in Figures 21, 22, and 23. The direction of elongation of the ion cloud is nearly perpendicular to the direction of motion of the ion cloud. The contours shown in Figure 21 show an ion cloud with a well-defined center and strong gradients. The contours of Figure 22 show a more elongated ion cloud that maintains strong gradients on its side. The third display of contours of Figure 23 shows a larger, but considerably less dense cloud with smaller gradients and a not-so-well defined shape.

We chose this later time to explore the vertical extent and shape of the ion cloud. Figure 24 shows the vertical profile of the ion cloud in the North-South vertical plane. According to Figure 22, the North-South dimension is close to the smallest dimension of the ion cloud, and certainly it is considerably smaller than the East-West dimension. Figure 24 also shows, for reference, the size of the radar beam and the area within which measurements were acquired. This area is within the limits marked as boundary for acquisition of data. The approximate direction of the Earth's magnetic field is also shown in this figure.

A few observations can be made at this point. The short vertical width of the ion cloud seems to be a real feature of this cloud: 9 km for 3-dB width. The overall shape of the ion cloud does not seem to align with the Earth's magnetic field. This last feature could be understood by saying that different regions of the ion cloud have descended to different altitudes. That is, if the northern part of the ion cloud which also happens to be the eastern part is 5 km lower than the southern part, then we can understand better the relation of the contours of Figure 24 to the magnetic field line. Again, these observations should be made with caution because of the limitations on the data acquired by the radar. The radar antenna beam uses a discrete number of positions to acquire data, and Figure 24 may be incomplete because not enough points were observed.

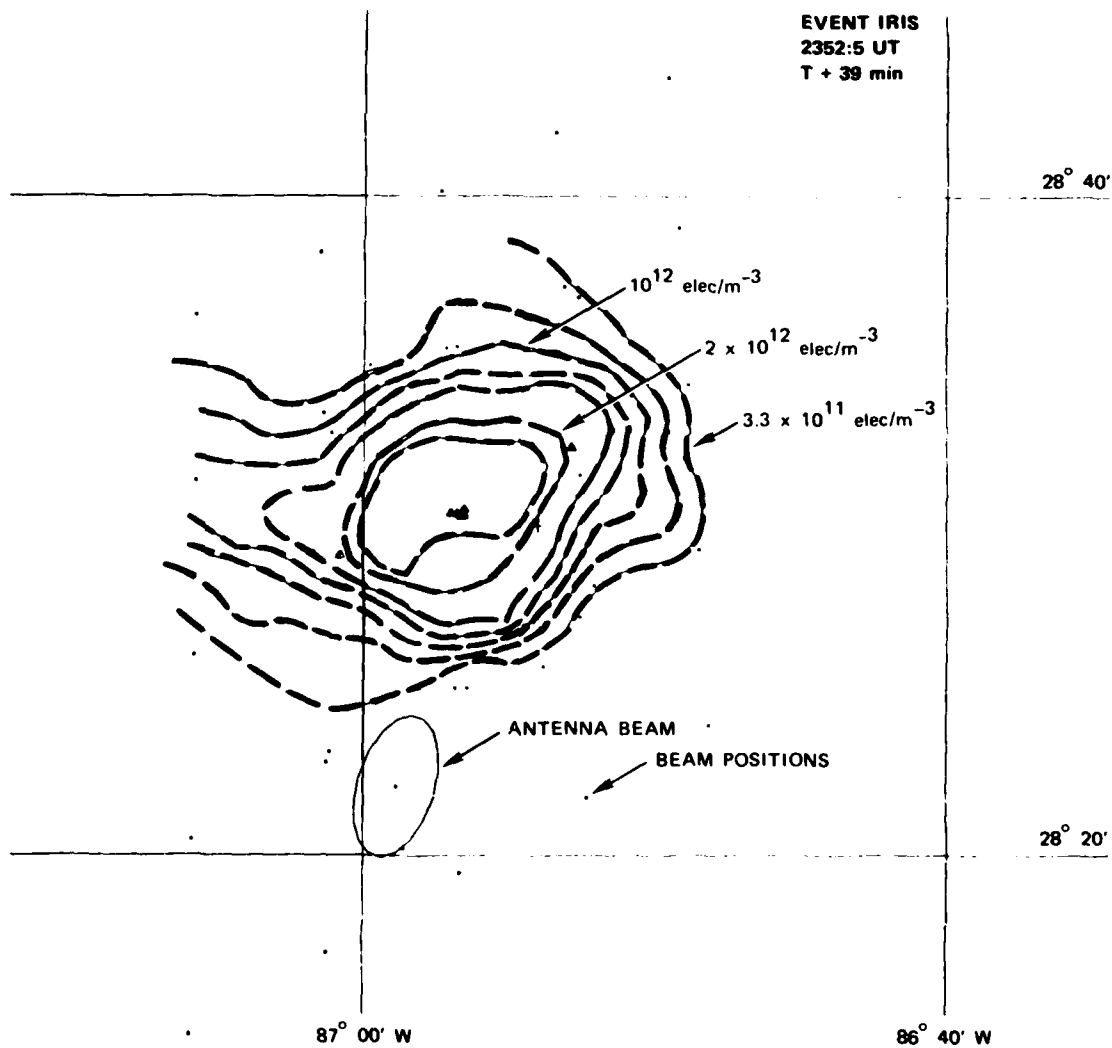


FIGURE 21 HORIZONTAL CONSTANT ELECTRON DENSITY CONTOURS AT T + 21 min

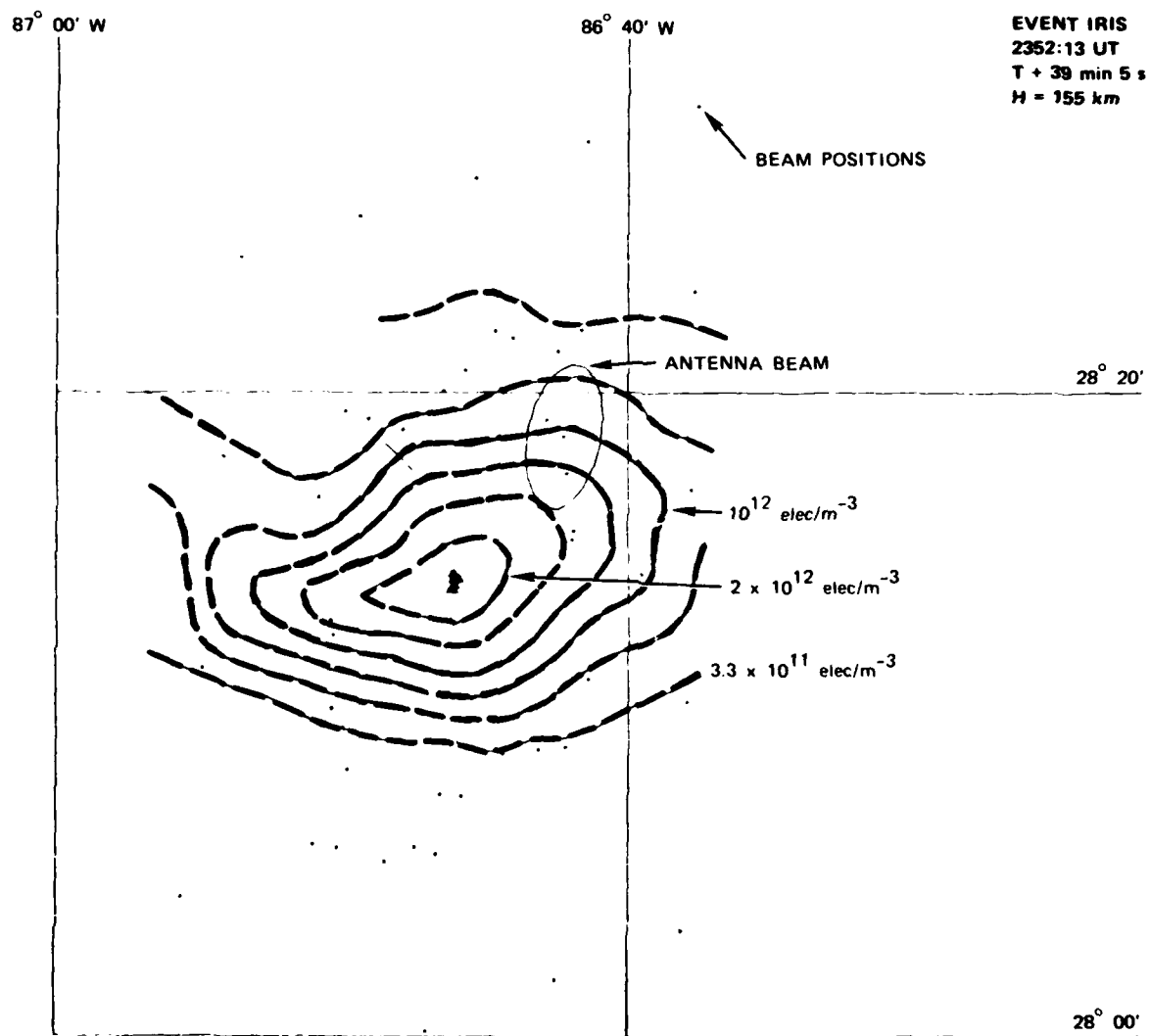


FIGURE 22 HORIZONTAL CONSTANT ELECTRON DENSITY CONTOURS OF EVENT IRIS
AT T + 39 min

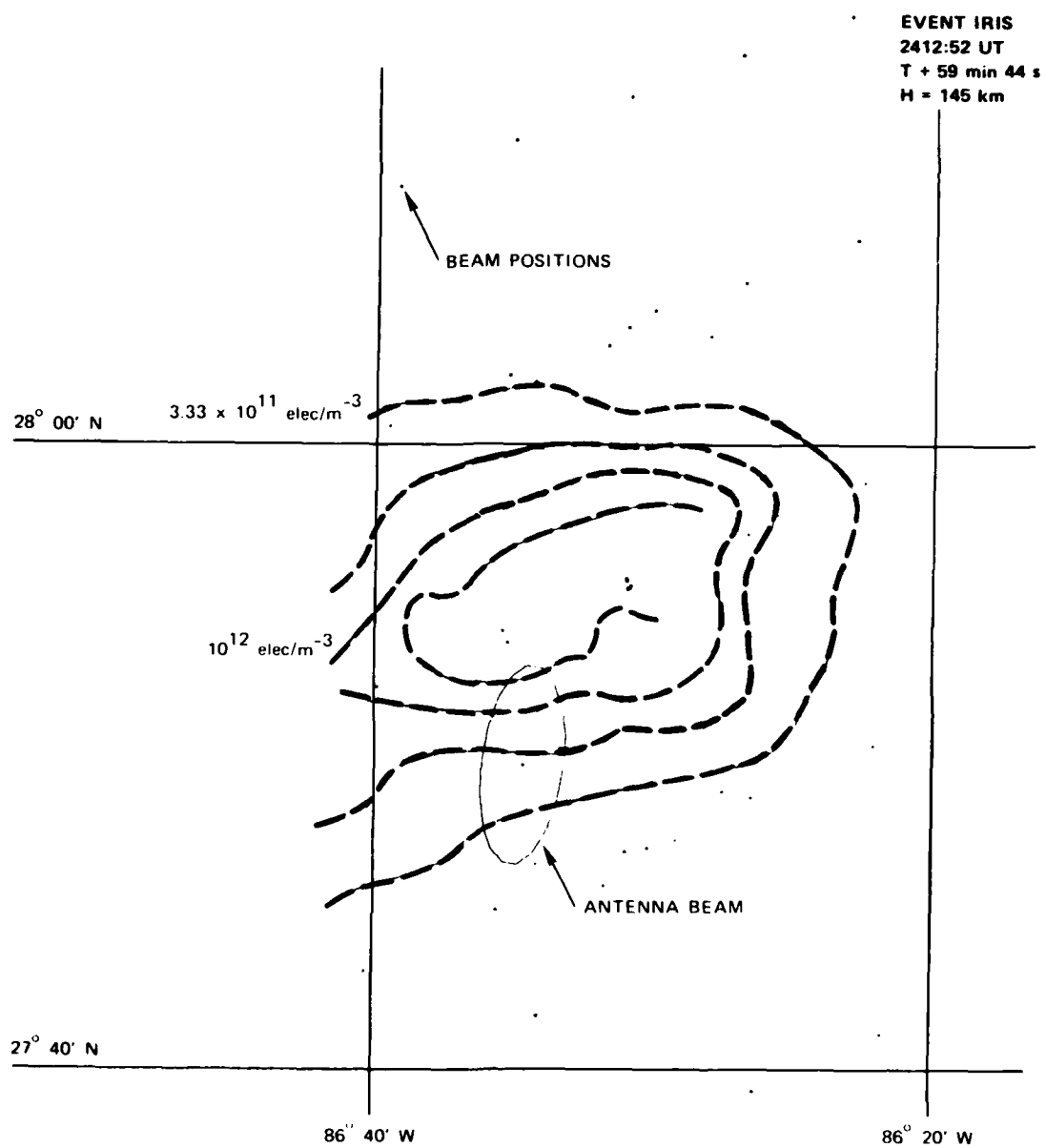


FIGURE 23 HORIZONTAL CONSTANT ELECTRON DENSITY CONTOURS OF EVENT IRIS
AT T + 60 min

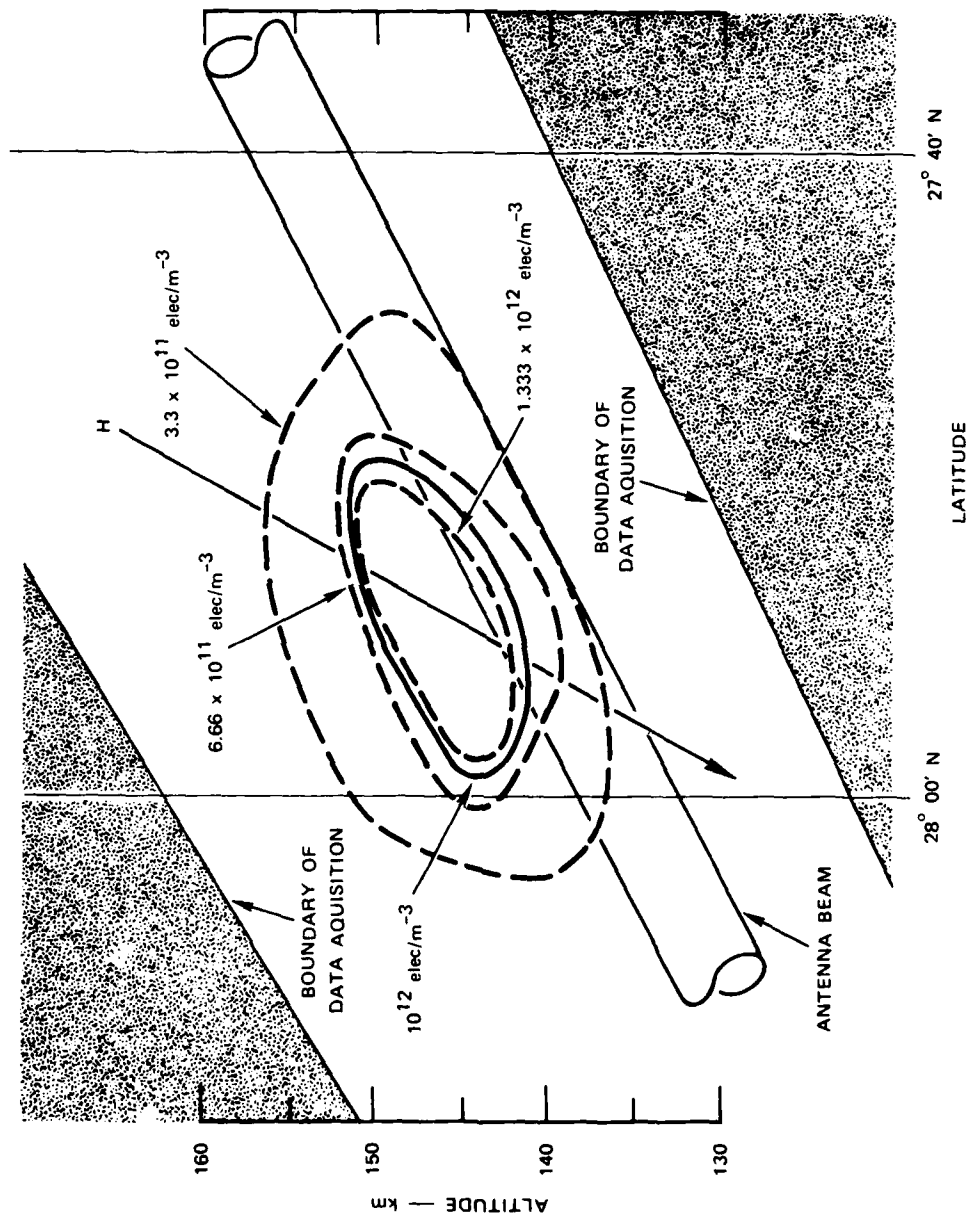


FIGURE 24 VERTICAL SLOPE OF EVENT IRIS AT 2413 UT (T + 60 min)

VI EVENT JAN

Event JAN was affected by two failures in the computer system. After a partial recovery between $T + 8$ min and $T + 25$ min, the second breakdown occurred and all efforts to reacquire the ion cloud were unsuccessful. The ground track, the altitude data, and the maximum electron density data obtained are presented in Figures 25, 26, and 27 for the sake of completeness. The very wide spread in the data points indicates the "no-track" situation that was present for most of the time.

The period of time between $T + 8$ min and $T + 25$ min may yield useful information for correlation with the probe rocket if the time for a careful data analysis is invested.

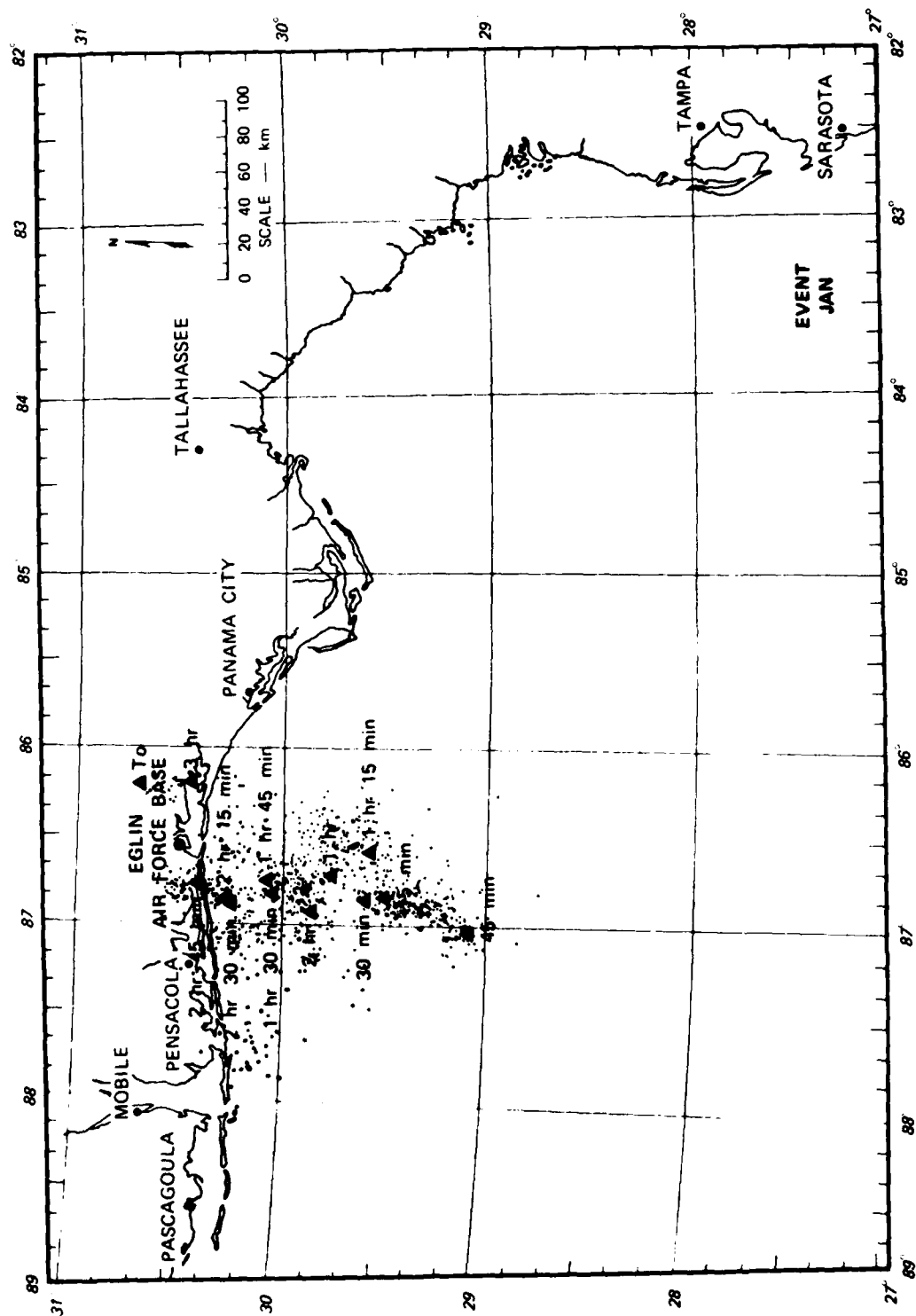


FIGURE 25 HORIZONTAL TRACK DATA OF EVENT JAN

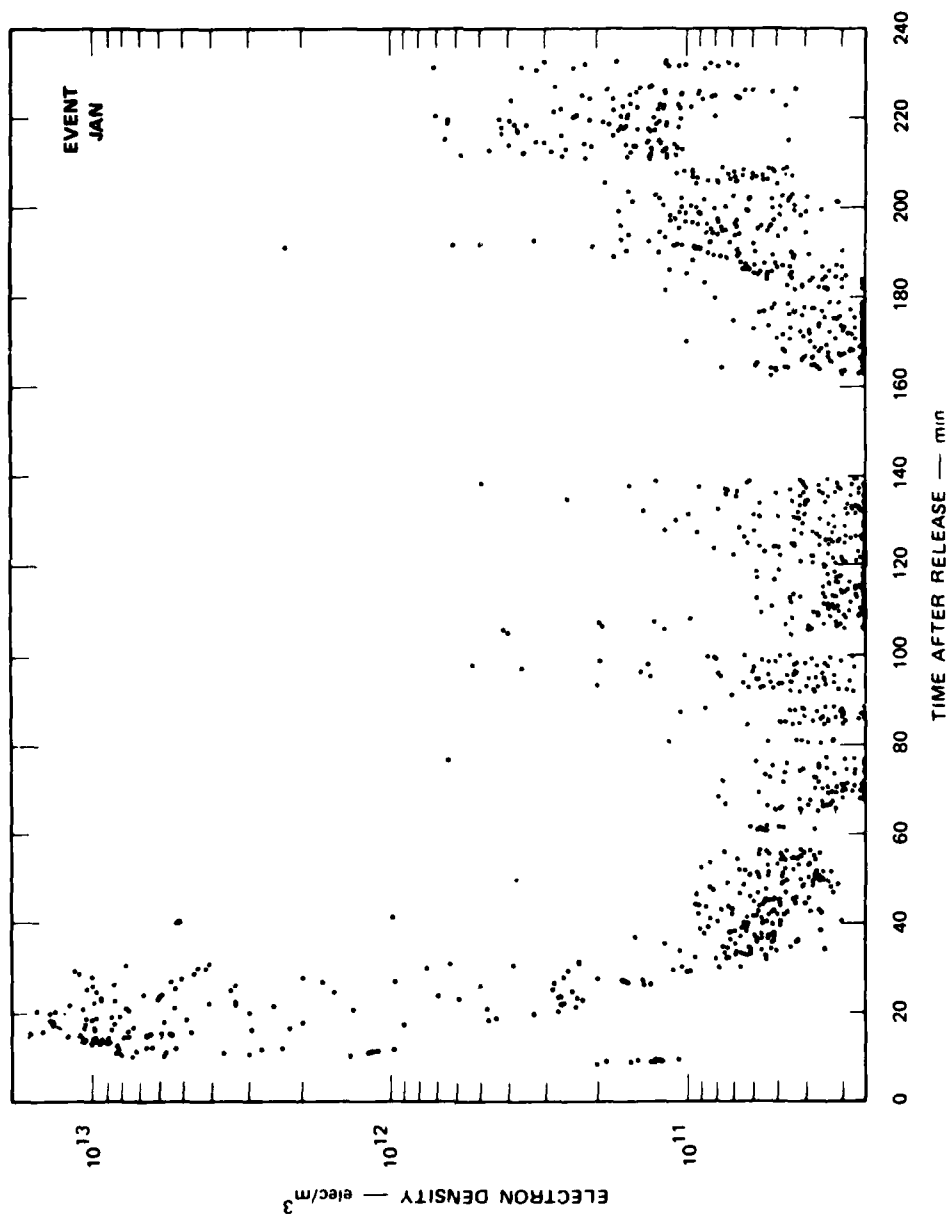


FIGURE 26 ELECTRON DENSITY DATA OF EVENT JAN AS A FUNCTION OF TIME

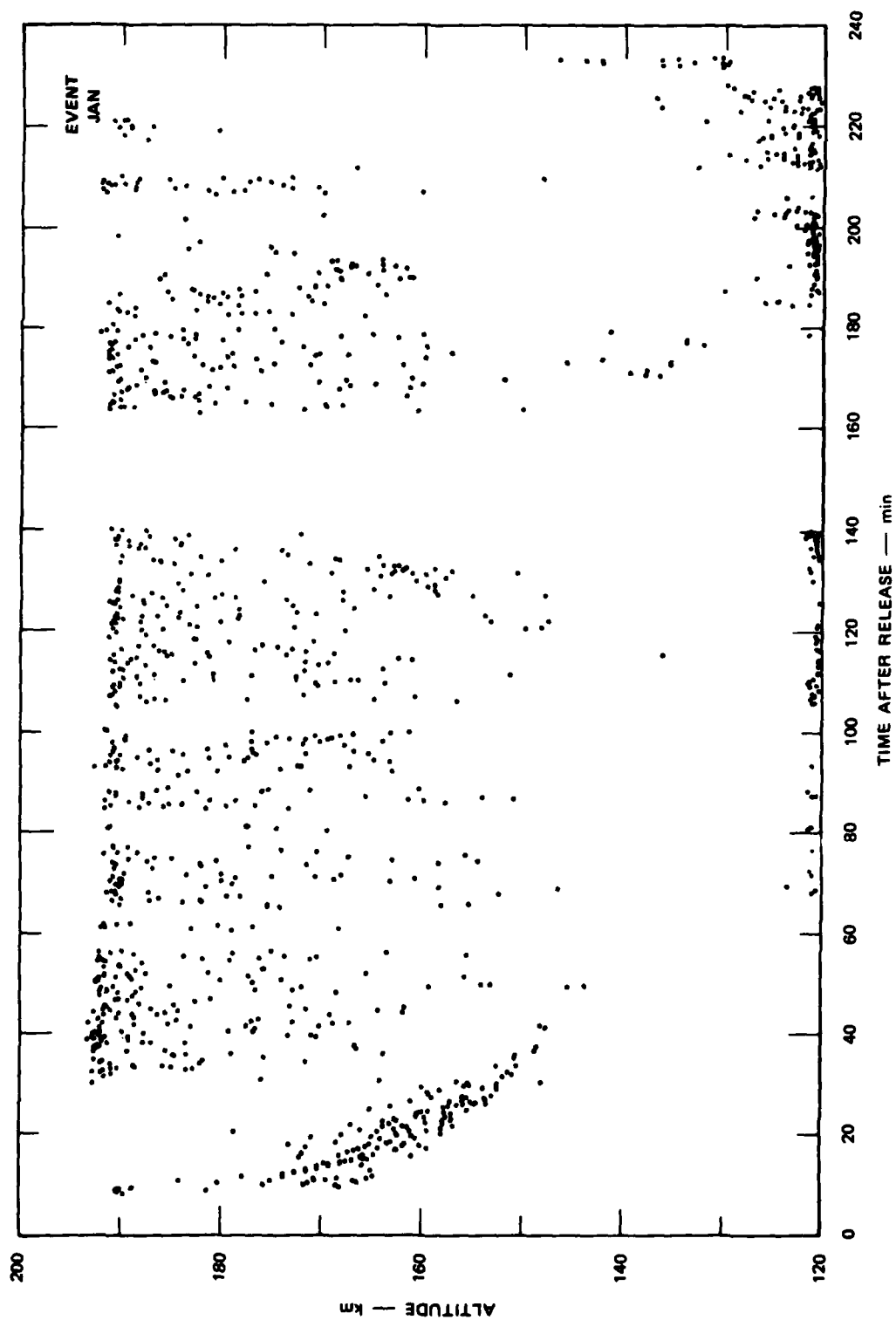


FIGURE 27 ALTITUDE DATA OF EVENT JAN AS A FUNCTION OF TIME

VII SUMMARY

During the PLACES series of experiments, Event GAIL was tracked for 2 hours, Event HOPE for 4 hours, Event IRIS for 1-1/2 hours; some data were obtained for Event JAN for about 15 min. The four-hour track of Event HOPE is the longest time a barium ion cloud has ever been tracked and the time could have been longer if we had not been constrained by the availability of the radar. Event JAN is singled out by the different extreme. It is the worst event in terms of tracking results. Thus, the PLACES series exceeded the STRESS series in one aspect and falls short of it in another.

Part of our shortcomings were caused by the aging and cumbersome computer system in the radar. The results of the programming are small compared with the effort applied. The new tracking algorithm, coupled with our new sampling hardware, did not start to work in time to be effective.

Were a new series of releases to take place in the future, several aspects of the work done could be picked up where we left off. The continuous change in the operating system of the FPS-85 will require some reprogramming. Event JAN has demonstrated the need for a more flexible tracking algorithm.

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